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Modeling of Atmospheric Structure, 70-130 km

GERALD V. GROVES



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consistency in nitrogen partial pressures between the 70 km and 130 km values of the given lower and upper models. This discrepancy which at present remains an unresolved problem is discussed in the text. Tables of temperature, pressure, and density are included in the report based on the best fit to available data and simultaneously satisfying the constraints of the upper and lower models.

The tables presented are of diurnal and zonal mean values of temperature, pressure and density at 5-km height intervals from 70 to 130 km for each mid-month date and solar activity of $F_{10.7} = 70$ and 150 units and geomagnetic activity of $A_p = 4$ and 132.

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Modeling of Atmospheric Structure, 70-130 km

1. INTRODUCTION

Season and latitude have traditionally been considered the major influence on the structure of the stratosphere and mesosphere. The Air Force Reference Atmospheres tabulated temperature, density and pressure for each month and each 15°N latitude to 90 km altitude. CIRA 1972, Part 2 tabulated these parameters for each month and each 10°N latitude to 110 km. Tabulations to 80 km at monthly intervals have been extended to the southern hemisphere with a 10° latitude interval (Koshelkov, Barnett, Barnett and Corney).

(Received for publication 15 June 1987)

^{1.} Cole, A. E., and Kantor, A. J. (1978) <u>Air Force Reference Atmospheres</u>, <u>Air Force Surveys in Geophysics</u>, <u>No. 382</u>, <u>AFGL-TR-78-0051</u>, <u>ADA 058505</u>, <u>Second Printing</u>, <u>Corrected edition</u>, <u>March 1984</u>.

^{2.} Koshelkov, Yu. P. (1983) Proposal for a reference model of the middle atmosphere of the southern hemisphere, Adv. Space Res. 3:3.

^{3.} Barnett, J. J. (1984) Plots and Tables of Temperature and Geopotential Height Based on Nimbus 5 SCR and Nimbus 6 PMR, Working Group 4 Document, XXV COSPAR Meeting, Graz, Austria.

^{4.} Barnett, J.J., and Corney, M. (1985) Middle atmosphere reference model derived from satellite data, in <u>Handbook for MAP</u>, 16:47, July 1985, SCOSTEP Secretariat, University of Illinois, 1406 W. Green Street, Urbana, IL 61801.

Additional dependences, notably local solar time and solar activity, are essential for modelling of the thermosphere as, for example, in CIRA 1972, Part 3 which extends upwards from 110 km. Further dependences are included in MSIS-83, 5 which extends upwards from 90 km.

The present report addresses itself to the problem of modelling the intermediate region between mesospheric and thermospheric models. It seeks to develop a procedure for generating a smooth transition between selected lower and upper models in all relevant neutral atmosphere parameters, including composition.

The need for improved modelling of the mesosphere/thermosphere region has long been evident and a 1984 recommendation for a new CIRA referred to the need for tabulations from 80 to 120 km. An altitude range from 70 to 130 km has been adopted for the present work to allow a smooth transition to be introduced to the lower and upper models, whose values are required to be matched at 70 and 130 km. The problem of modelling the 70 to 130 km region is therefore one of satisfying two types of conditions. One is imposed by the requirements of the upper and lower models and the other by the temperature data of the intermediate region. Such data are taken in the present work at 5 km height intervals, that is, at 75, 80, ..., 125 km. The problem is a rovel one in atmospheric modelling in that the usual requirement is of the latter type, that is, to obtain a fit to a set of data. In the present problem, the requirement to also match lower and upper models in all their relevant parameters is not an independent one, and a particular matter of interest and concern is the extent to which both types of condition can be satisfactorily met. The treatment here is to rigidly match the lower and upper models (with continuity in the second height derivative) and at the same time optimize the fit to the intermediate temperature data, which may show a residual bias due to the rigid end constraints. In previous works by Forbes and Groves, 6 the temperature data were found to be biased higher than model values and the matter warranted further consideration.

The procedure of model formulation is described in Appendix A and relates to the choice of MSIS-86 7 for the upper model and to that described in 'A Global Reference Atmosphere From 18 to 80 km' 8 for the lower model. A report of the

^{5.} Hedin, A. (1983) A revised thermospheric model based on mass spectrometer and incoherent scatter data, MSIS-83, J. Geophys. Res. 88:10170.

^{6.} Forbes, J.M., and Groves, G.V. (1986) Atmospheric Structure between 80 and 120 km, presented to Workshop No. XV, XXVI COSPAR Meeting, Toulouse, France.

^{7.} Hedin, A. (1986) CIRA 1986 Atmospheric Model in the Region 90 to 2000 km, Draft of 13 June 1986.

^{8.} Groves, G.V. (1985) A Global Reference Atmosphere From 18 to 80 km, Air Force Surveys in Geophysics, No. 448, AFGL-TR-85-0129, ADA 162499.

earlier calculations undertaken at a preliminary stage of this work with MSIS-83 as the upper model has been given previously. The subsequent stages then proposed for the formulation have now been completed and are reported here. These include (a) modelling of individual constituent gases and hence of total density and pressure, (b) combining the three models for the height regions 18 to 70, 70 to 130 and above 130 km to provide a single model from 18 km upwards (Appendix D, Section 3), and (c) the development of models for various solar conditions (and not only mean solar conditions as treated earlier.)

2. TEMPERATURE DATA UTILIZED, 70 TO 130 km

Temperature data have been assembled from two earlier collations of data, that is, from that used for the construction of CIRA 1972, Part 2 and the rocket and incoherent scatter (I.S.) data reviewed by Forbes. Analyses of I.S. data from St. Santin have also been utilized.

I.S. temperatures were utilized as monthly means for the locations and heights shown in Table 1 and were compared with models computed for the solar activity parameters shown in the table.

Table 1. Distribution of Available Monthly Mean I.S. Temperatures

	Height	Solar Activit	y
Site	km	$F_{10.7} = F_{10.7}$	Ap
Millstone Hill (42 N) (Wand, 11 Table 1)	105-125	120	7
St. Santin (45 N) (Alcayde et al, 10 Table 1)	95-110, 120	70, 120, 170	7
Arecibo (18 N) (1970-1975 data, Forbes ⁹)	105-130	108	10

^{9.} Forbes, J. M. (1984) Temperature Structure of the 80 km to 120 km Region, presented at the XXV COSPAR Meeting, Graz, Austria.

^{10.} Alcaydé, D. et al (1979) Temperature, molecular nitrogen concentration and turbulence in the lower thermosphere inferred from incoherent scatter data, Ann. Geophys. 35:41.

^{11.} Wand, R.H. (1983) Lower thermospheric structure from Millstone Hill incoherent scatter radar measurements 2. Semidiurnal temperature component, J. Geophys. Res. 88:7211.

Monthly mean temperatures from rocket techniques were utilized for the locations and heights shown in Table 2. The Kwajalein data of 1976-1978 have been compared with models computed for $F_{10.7} = \overline{F}_{10.7} = 95$ units. For the other data, which extend no higher than 100 km, solar effects are minimal and a value of $F_{10.7} = \overline{F}_{10.7} = 120$ units has been adopted for computing models for comparison purposes. Likewise a mean value of A_p (= 10) has been adopted in model computations.

Table 2. Distribution of Monthly Mean Temperatures Obtained by Rocket Techniques

Site	Height km	Site	Height km
Heiss Is. (81 N)	75-100	Thumba (8 N)	75
Volgograd (48 N)	75-100	Ships (equator)	75-100
Kwajalein (9 N)	75-120	Molodezhnaya (68 N)	75-80

The locations at which single rocket profiles of temperature were utilized are shown in Table 3. The data available fall off rapidly above 100 km as shown in Table 3 by the numbers of profiles available at 105 and 120 km. Single profiles are compared with models computed for the same solar activity parameters as defined by the procedure described in Appendix B.

Table 3. Numbers of Available Temperature Data at 105 and 120 km From Rocket Launchings

Site		Height 105	(km) 120	Site		Height 105	(km) 120	Site	Height 105	(km) 120
Pt Barrow	(71N)	3	0	Wallops	(38 N)	18	16	Barking S (22N)	3	0
Churchill	(59N)	8	5	White S	(32 N)	8	0	Kwajalein (9 N)	3	2
Sardinia	(44 N)	0	1	Eglin	(CO N)	6	1	Ascension (8S)	2	0

Developed into a monthly model (Cole et al 12) and listed in Table 2.

^{12.} Cole, A.E., Kantor, A.J., and Philbrick, C.R. (1979) <u>Kwajalein Reference Atmospheres</u>, 1979. <u>Environmental Research Papers</u>, No. 677. AFGL-TR-79-0241, ADA 081780.

3. MODEL FORMULATION, 70 TO 130 km

The procedure devised for calculating a model temperature profile from 70 to 130 km is set out, step-by-step, in Appendix A. Theoretical details are presented in Appendixes A1 to A5. An essential feature of the method is the expression, for given geophysical conditions, of g/T (g being gravity acceleration and T temperature) as a polynomial in height whose coefficients are chosen (a) to give continuity up to the second height derivative with the lower and upper models, (b) to reproduce, on integration of appropriate physical equations, the required ratio of N_2 pressure at 70 km to that at 130 km, and (c) to produce a best fit to observed temperatures at 75, 80, ..., 125 km altitude.

Appendix A sets out the step-by-step procedure for calculating density. The method is based on the MSIS-86⁷ formulation for number densities of individual gas constituents. Other parameters then follow as set out in Appendix A.

4. COMPARISON OF OBSERVED TEMPERATURES WITH MODELS

This section analyzes the differences between observed temperatures and model values computed in each case for the same geophysical conditions. Model values are derived by the procedure of Appendix A which involves polynomial coefficients as whose determination depends on the analysis of temperature data as described in Appendix B.

Three cases with different selections of temperature data have been considered and a set of coefficients $a_{\hbox{sn}}$ corresponding to the second of these cases is listed in Appendix B.

4.1 Case 1: a_{sn} Based on All Data Excluding Millstone Hill

Case 1 was undertaken in the preliminary investigation⁶ which utilized all data at first and then excluded Millstone Hill data because of their relatively high values. The comparisons reported in Sections 4.1.1 to 4.1.3 essentially repeat and update the results previously given.

4.1.1 a_{sn} AT LOW LATITUDES BASED ON ALL DATA

The first two columns of Table 4 shows the average deviations from the computed models of temperature data for all months at Kwajalein (9N) and Arecibo (18 N). Such deviations are denoted by $\mathbf{x_i}$ (for the i th site) in Appendix C which presents the relations used in the analysis. Kwajalein and Arecibo provide the main input of data

above 100 km at low latitude and as previously noted by Forbes and Groves, ⁶ the Kwajalein temperatures are lower than those of Arecibo by about 15 K on average.

Table 4. Low Latitude Temperature Differences From Computed Model Values (K). Av = average difference for data of all months at a particular site. Case 1

Height	leight Kwajalein		Arec	Arecibo		Ships		sites
km	Av	sd	Av	sd	(equa Av	tor) sd	Mean of Av	sd
75	-1.4	1.3	-	-	-0.5	1.6	-0.5	1.7
8 0	2.4	0.8	-	-	-3.3	1.2	1.4	1.5
85	5.9	0.8	-	-	-6.5	1.0	1.8	3.2
90	6.9	0.7	-	-	-6.0	1.2	3.8	2.8
95	3.9	0.8	-	-	-1.7	2.1	3.3	1.7
100	-4.4	0.9	-	~	9.3	4.6	-3.5	2.6
105	-8.0	1.8	9.9	5.2	-	-	-3.5	13.0
110	1.4	1.3	13.9	5.6	-	_	2.1	3.5
115	10.3	3.9	23.0	6.1	-	-	14.0	8.2
120	-9.8	2.7	7.6	4.6	-	-	-5.3	10.7
125	-	-	-10.8	7.0	-	-	-10.8	7.0**

^{*}Ascension Is. (8 S), Natal (6 S), Ships (equator), Kourou (5 N), Thumba (8 N), Kwajalein (9 N), Arecibo (18 N), Barking Sands (22 N), Carnarvon (25 S).

An unsatisfactory latitude variation at low latitudes was modeled in the preliminary work by closely fitting to these data. The difficulty has been overcome in the present work by increasing the degree of smoothing to give models with an acceptable latitude structure at low latitude. The lower Kwajalein temperatures then lead to lower differences as shown for 105 to 120 km in Table 4.

In the last column of Table 4 are the mean values $(\bar{\mathbf{x}})$, where the mean is taken over all sites appropriately weighted. The 11 means at 75, ..., 125 km are both positive and negative with a weighted average of about 1 K. A 1 K change over the height range 70 to 130 km is equivalent to a 4 percent change in the ratio of N_2 pressures at 70 and 130 km. Such changes would be within the expected limits of accuracy and therefore the intermediate temperature data, taken as a whole over all sites, are not inconsistent with the lower and upper models. Data for a particular site may nevertheless still be biased one way or the other with respect to

^{**}Arecibo data only.

the models, for example, Arecibo data at 105 to 120 km show a region of bias to higher temperatures by about 10 to 20 K.

4.1.2 a_{sn} AT MIDDLE LATITUDES BASED ON ALL DATA EXCLUDING MILLSTONE HILL

Table 5 presents average temperature differences from computed models for the five middle latitude sites that contribute most data at heights above 100 km. The mean of such averages with respect to all sites from 30° to 50° latitude is also shown. In contrast to the results of Table 4, model values are consistently lower than observed temperatures. Such a bias was noted and reported at the preliminary stage of this work and is able to arise as the least-squares fit is constrained by the requirement to match the N_2 pressures of the lower and upper models at 70 and 130 km.

Table 5. Middle Latitude Temperature Differences From Computed Models (K). Av=average difference for data of all months at a particular site. Case 1

Height	Egl	lin		hite inds	Wal	lops	St San		Mill: Hi	stone 11	All sit Mean of	
km	Av	sd	Av	sd	Av	sd	Αv	sd	Αv	sd	Av	sd
75	-2	2	9	4	2	1	-	-	-	_	0.3	1.1
80	1	4	9	4	0	2	-	-	-	-	1. 1	0.9
85	2	3	3	5	3	2	-	-	-	-	5.8	1.7
90	11	3	4	4	8	2	-	-	-	-	10, 2	1.9
95	20	6	15	5	14	4	8	1	-	-	9, 8	1.5
100	24	7	21	7	13	5	1	1	-	-	6, 1	5.1
105	24	4	27	8	5	5	3	1	22	1	13, 2	6.2
110	46	19	14	8	7	10	10	2	38	3	24, 5	12.0
115	-	-	-	-	54	14	-	-	49	3	48.9	1.6
120	-	-	-	-	48	18	32	3	39	3	35,9	4.6
125	-	-	-	-	27	25	-	-	13	2	12.8	1.5

^{*}Eglin (30 N), Woomera (31 S), White Sands (32 N), Arenosillo (37 N), Wallops (38 N), Millstone Hill (42 N), Sardinia (44 N), St. Santin (45 N), Volgograd (48 N), Kerguelen (49 S).

If we consider the changes that would be needed in the models at 70 or 130 km to enable the intermediate model temperatures to fit observed values we find that either (a) pressures at 70 km would need to be lower by 40 percent or (b) pressures at 130 km would need to be higher by 40 percent or (c) corresponding partial adjustments would need to be made at both heights. Such pressure adjustments would require lower temperatures over some range of height in either or both of the lower and upper models thereby allowing higher values to be modeled in the intermediate region. The discrepancy between observed and model values is discussed in Section 7.

4.1.3 a_{sn} AT HIGH LATITUDES BASED ON ALL DATA

Table 6 shows average temperature differences from computed models for three high latitude sites. Data are available from few sites at high latitude and then mainly below 105 km. As would be expected in these circumstances, no difficulty arises in deriving models that are consistent with the limited height range of available data.

Table 6. High Latitude Temperature Differences From Computed Models (K). Av = average difference for data of all months at a particular site. Case 1

Height	Churchill		Pt. Barrow		Heiss	Is.	All sites*		
km	Av	sd	Av	sd	Av	sd	Mean of Av	sd	
75	3.6	1.6	-1.4	2.2	0.3	1.3	-1.4	3.1	
80	-0.2	1.9	-0.4	2.4	0.6	1.8	-1.1	1.4	
85	0.4	1.7	-0.2	3.0	0.7	1.6	0.5	0.4	
90	5.8	3.1	6.7	2.8	-1.7	2.0	2.1	4.9	
95	11.2	4.6	15.7	8.3	-5.8	2.7	-0.2	10.3	
100	14.2	4.1	-7.6	13.1	-13.2	2.8	-4.6	15.4	
105	12.0	7.4	-24.8	3.6	-	-	-17.8	20.5	
110	6.0	13.3	-	-	-	-	6.0	13.3**	
115	26.3	19.7	-	-	-	-	26.3	19.7**	
120	47.3	32.3	-	-	-	-	47.3	32. 3**	
125	-	-	-	-	-	-	_	_	

^{*}Churchill (59 N), Molodezhnaya (68 S), Pt. Barrow (71 N), Heiss Is. (81 N).

^{**}Churchill data only.

Table 6 shows the mean of the average differences with respect to the few existing high latitude sites. The absence among these means of any significant difference from zero is a result of the paucity of data above 100 km.

4.2 Case 2: a_{sn} Determined Without I.S. Data and Without Data From 105 to 125 km

By excluding data above 100 km from the analysis of Appendix B and the determination of $a_{\rm Sn}$, we can expect to compute models that have a better fit to data at and below 100 km at the expense of having a poorer fit to the excluded data above 100 km. Comparisons between observed and computed temperatures are presented for the same groups of low, middle and high latitude sites as for Case 1. Data are much more numerous below 100 km than above and for this reason Case 2 is worthy of investigation.

4.2.1 a AT LOW LATITUDES DETERMINED WITHOUT I.S. DATA AND WITHOUT DATA FROM 105-125 km

Table 7 presents the results for Case 2 corresponding to those for Case 1 in Table 4. As would be expected, model temperatures at 110 - 125 km are now lower (by ~ 10 K) and those at 85 - 100 km are higher (by ~ 2 K). These changes are generally significant for any particular site but over all sites the last columns of Tables 4 and 7 show that the changes in the means (which are also ~ 10 K and 2K in the respective height ranges) are not significant, being of the same order of magnitude as the standard deviations (sd) of the means.

The fact that omission of data above 100 km results in such a limited change, on average, in the modeling accords with the conclusion of Section 4.1.1 that observed temperatures for 75 - 125 km are not inconsistent with the lower and upper models when taken as a whole over all sites.

For individual sites, the average differences between observed and computed temperatures may be significantly increased or decreased for Case 2 relative to Case 1. For Arecibo, observed values at 110 - 120 km are now higher than computed values by 30 - 40 K.

Table 7. Low Latitude Temperature Differences From Computed Model Values (K). Av = average difference for data of all months at a particular site. Case 2

Height			Arec	ibo	Shir (equa		All site Mean of	s *
km_			Av sd		Av	sd	Av	sd
75	-1.3	1.3	•	-	-0.5	1.6	-0.4	1.7
80	2.5	0.8	-	-	-3.7	1.1	1.1	1.6
85	4.8	0.7	-	-	-7.5	0.9	0.4	3.2
90	4.2	0.5	-	-	-7.8	1.0	1.7	2.5
95	0.4	0.8	-	-	-3.6	1.8	-0.1	1.4
100	-7.1	1.0	-	_	8.1	4.3	-ô. 2	2.8
105	-8.0	1.9	9.3	5.7	-	-	-2.8	14.9
110	5.4	1.3	23.3	5.8	-	-	6.4	4.9
115	18.6	4.0	42.3	6.2	-	-	25.6	15.3
120	-0.7	2.7	29.9	5.3	-	-	5.6	17.5
125		-	-2.1	6.8	-	-	-2.1	6.8**

^{*}Ascension Is. (8 S), Natal (6 S), Ships (equator), Kourou (5 N), Thumba (8 N), Kwajalein (9 N), Arecibo (18 N), Barking Sands (22 N), Carnarvon (25 S)

4.2.2 a_{sn} AT MIDDLE LATITUDES DETERMINED WITHOUT I.S. DATA AND WITHOUT DATA FROM 105-125 km

Table 8 presents average differences for Case 2 corresponding to Table 5 for Case 1. As expected, the fit up to 100 km is much improved with both positive and negative average differences whose means over all sites are now not significantly different from zero (the last column of Table 8).

Above 100 km the high average differences of Table 5 become still higher for Case 2 being $\sim 80 \text{ K}$ at 115 - 120 km.

At middle latitudes, Cases 1 and 2 present straight choices between models that are biased lower than observations at all heights, 75 - 125 km, (Case 1) and those that are biased lower than observations at only 105 - 125 km (Case 2), the biases nevertheless being significantly greater.

^{**} A recibo data only.

Table 8. Middle Latitude Temperature Differences From Computed Models (K). Av = average difference for data of all months at a particular site. Case 2

Height	Eglin		White Sands		Wallops		St. Santin		Millstone Hill		All sites* Mean of	
km	Av	sd	Av	sd	Av	sd	Av	sd	Av	sd	Av	sd
75	- 1	2	10	4	3	1		-	-	-	1.0	1.0
80	2	4	9	4	1	1	-	~	-	-	2.0	0.9
85	-2	3	- 1	5	0	2	-	~	-	-	2. 1	1.8
90	-1	3	-6	4	-3	2	-	~	-	-	-0.3	1.6
95	3	5	2	4	- 1	4	-6	1	-	-	-2.9	2.0
100	11	7	11	6	1	5	-11	1	-	-	-5.7	5.3
105	23	4	26	8	4	5	1	1	21	1	13.8	6.0
110	6 2	19	28	9	25	10	23	2	53	2	42.4	12.8
115	_	-	-	-	91	14	-	-	80	3	80.8	2.9
120	-	-	-	-	90	17	65	3	75	3	69.9	6.3
125	-	-	-	-	45	25	-	~	27	2	27.2	1.7

^{*}Eglin (30 N), Woomera (31 S), White Sands (32 N), Arenosillo (37 N), Wallops (39 N), Millstone Hill (42 N), Sardinia (44 N), St. Santin (45 N), Volgograd (48 N), Kerguelen (49 S)

4.2.3 a AT HIGH LATITUDES DETERMINED WITHOUT I.S. DATA AND WITHOUT DATA FROM 105-125 km

Table 9 shows average differences at high latitude sites when data at 105 km and above are omitted from the determination of $a_{\rm sn}$. At 110 km and above the omitted data amount to only eight rocket temperature profiles at Churchill (59 N) and hence Table 9 shows no significant change from Table 6, being given here for the sake of completeness.

Table 9. High Latitude Temperature Differences From Computed Models (K). Av = average difference for data of all months at a particular site. Case 2

Height	Churchill		Pt. Barrow		Heiss	Is.	All S Mean of	Sites*
km	Av	sd	Av	sd	Av	sd	Av	sd
75	4.3	1.6	-1.2	2.2	0.2	1.3	-1.2	3.2
80	1.5	1.9	0.3	2.4	0.4	1.8	-0.3	1.4
85	0.4	1.7	0.0	3.0	0.6	1.6	0.4	0.2
90	1.9	3.0	5.9	2.8	-1.6	2.0	1.2	3.9
95	3.9	4.7	14.2	8.2	-5.3	2.7	-1.7	7.3
100	7.5	4.2	-9.3	13.1	-12.7	2.8	-6.6	11.3
105	9.4	7.6	-25.7	3.6	-	-	-19.3	19.2
110	10.8	13.6	-	-	-	-	10.8	13.6**
115	40.0	19.9	-	-	-	-	40.0	19.9**
120	64.3	32.1	-	-	-	-	64.3	32.1**
125	_			_		-	-	-

^{*}Churchill (59 N), Molodezhnaya (68 S), Pt. Barrow (71 N), Heiss Is. (81 N).

4.3 Case 3: a Determined Without Data at 75 to 95 km

In this case, all data at heights 100 - 125 km, including incoherent scatter data from Arecibo, Millstone Hill and St. Santin have been utilized for determining a according to the procedure of Appendix B, while that below 100 km have been omitted. In comparison with Cases 1 and 2 an improvement in the fit of computed temperatures to observed values can be expected above 100 km at the expense of lower computed temperatures and a poorer fit to observations below 100 km.

Table 10 shows the results obtained for mid-latitude data corresponding to Tables 4 and 7 for Cases 1 and 2. At 85 - 95 km, Table 10 shows the temperatures are ~ 20 K higher than model values. In comparison, the difference is ~ 10 K for Case 1, but differences above 100 km are greater. Case 2 produces a close fit at 85 - 95 km with still larger differences above 100 km.

^{**}Churchill data only.

Table 10. Middle Latitude Temperature Differences From Computed Models (K). Av = average difference for data of all months at a particular site. Case 3

Height	Εg	glin	in White Sands		Wallops		St. Santin		Millstone Hill		All Sites* Mean of	
km	Av	sd	Av	sd	Av	sd	Av	sd	Av	sd	Av	sd
75	0	2	12	4	5	1	-	_	-	-	2.6	1. 2
80	10	4	20	4	11	2	-	-	-	-	10.8	1.0
85	16	3	19	5	19	2	-	-	-	-	20.3	1.7
90	23	3	18	4	23	2	-	-	-	-	23.4	1.9
95	25	5	21	5	22	4	16	1	-	-	18.5	1.6
100	19	7	17	7	10	5	- 1	1	-	-	4.4	5.3
105	10	4	10	8	- 12	5	-12	1	6	2	-6, 5	5.7
110	23	18	- 14	8	-23	9	- 19	2	8	4	-2,4	14. 1
115	-	-	-	-	19	12	-	-	13	4	13,3	2.6
120	-	-	-	-	23	16	6	4	13	4	9,9	5.3
125	-	-	-	-	22	25		_	6	1	5, 9	1.3

^{*}Eglin (30N), Woomera (31S), White Sands (32N), Arenosillo (37N), Wallops (38N), Millstone Hill (42N), Sardinia (44N), St. Santin (45N), Volgograd (48N), Kerguelen (49S)

5. TABULATIONS OF TEMPERATURE, PRESSURE AND DENSITY

A wide choice of options is open for the particular models to present as tables and for the format in which they are to appear. The choice involves the height interval of tabulation and many geophysical parameters.

Diurnal and zonal mean values have been tabulated in the course of this project as latitudinal-height cross-sections for solar activity parameters $F_{10.7}$ = 70 and 150 units and A_D = 4 and 132.

Two formats have been adopted and coded (Appendix D, Section 3). By taking a 5-km height interval (70-130 km) and 20° latitude interval (80 S to 80 N) in the program TVAL5KMP, 12 tables consisting of temperature, pressure and density cross-sections for four selected months are generated that fit together on one page. Such tables are presented in Appendix F for all months based on Case 1 determinations of a_{sn} and four cases of solar activity: $F_{10.7} = 70$, 150 units, $A_p = 4$, 132.

By taking a 1-km height interval and 10° latitude interval from 80S to 80 N in program TVAL1KMP, tables are generated with a format identical with that previously used (Groves⁸) for 18-80 km with each table occupying one page. Both programs generate tables of W-E geostrophic wind.

6. LIMITATIONS OF THE MODEL FORMULATION, 70 TO 130 km

6.1 Types of Data

The only type of data that has been utilized at heights 75 - 125 km is temperature, which has served both as an input to the formulation of the models for these heights and for comparison with the computed models to check their goodness-of-fit. A computed composition is available from the formulation, but density data have not been utilized either as an input to the formulation or to check against computed values. No check has been made between W-E winds computed from the geostrophic wind equation and observed values.

6.2 Local Solar Time

No representation of tidal components at 70-130 km is included in the formulation. At 130 km, dependence on L.S.T. is that defined by MSIS-86 while, at 70 km, the lower model is independent of L.S.T. At intermediate heights the L.S.T. dependence arises by interpolation while ignoring tidal fields that may have detailed spatial structures.

6.3 Longitude and Solar Activity Effects

No detailed representation of longitudinal variations at 70 - 130 km is included in the formulation. At 70 and 130 km, the dependences on longitude are those of the adjacent lower and upper models and, relative to other variations at these heights, are very small. At intermediate heights, longitude dependence is then generated by interpolation and is a correspondingly small effect.

Dependence on solar activity is also not directly represented at heights between 70 and 130 km, but is present in the interpolation between these two heights where it is represented.

6.4 The Coefficients a sn

Sets of polynomial coefficients a_{sn} (s = 1, ..., S; n = 1, ..., N) are introduced by which height-latitude cross-sections of atmospheric structure parameters may be generated from 70 to 130 km at all latitudes. Subscript n is associated with a polynomial in height and subscript s with a polynomial in sin(latitude). Lack of data limits the number of sets of coefficients determined to 12, one for each month of the year. In principle, additional sets could be introduced with dependences on L.S.T., longitude, and solar activity to help overcome the limitations mentioned in Sections 6.2 and 6.3, but a vastly greater amount of data than that available would be required.

After many test computations with different choices of N and S, the choice was made of N=2 and S=7, giving 14 coefficients per set. Without improvements in the quality and quantity of available temperature profiles and their latitude distribution, N and S could not be justifiably increased.

Since a_{sn} relate to a pole-to-pole representation, they provide a means of formulating seasonal asymmetries between N and S hemispheres. Unfortunately observations are not available for much of the S hemisphere and it becomes necessary to assume seasonal symmetry for latitudes where observations are lacking in the determination of a_{sn} (Appendix B).

7. DISCUSSION

The modeling of atmospheric properties in an intermediate height region (70 - 130 km) between given mesospheric and thermospheric models has been treated theoretically (Appendixes A, A1 to A5, and B) and computationally as summarized in Appendixes D, E and F.

Temperature data (at 75, 80, ..., 125 km) that are utilized for modeling the intermediate region are summarized in Section 2 and two formats that have been devised for tabulating models of temperature, pressure and density, are outlined in Section 5 with examples given in Appendix F.

Section 4 presents comparisons between observed temperatures (75 - 125 km) and model values and is a section that commands particular attention on account of biases between the two that point to a possible inconsistency between the observed temperatures (75 - 125 km) and the given lower and upper models that are matched at 70 and 130 km. The biases feature strongly at mid-latitudes, where the majority of available data have been obtained, and are in the sense that observed values are higher than model values (Case 1 and Table 5).

One way in which the discrepancy could be resolved would be by modifying the $\rm N_2$ pressures of the lower and upper models at one or both of the heights 70 and 130 km. The magnitude of the discrepancy is such that its removal would require a 40 percent change in the $\rm N_2$ pressure ratio for these two heights and, as only a few percent adjustment could reasonably be allowed at 70 km, the majority of it would need to be applied at 130 km.

Another possible conclusion is that observed temperatures are erroneously high over at least some part of the 70-130 km height region. Above 100 km, temperature data are provided by the I.S. technique and a small number of measurements by rocket techniques as summarized in Tables 1 to 3. Up to 100 km, rocket data are much more numerous and are obtained by well-established techniques, such as the grenade experiment. Case 2 (Section 4.2) was therefore introduced to provide models that are consistent with the data at least up to 100 km (Table 8). However, the closer fit up to 100 km is obtained at the expense of a poorer fit above 100 km as shown in Table 8, where biases at 115-120 km are close to 80K for both I.S. measurements at Millstone Hill and St. Santin as well as rocket techniques at Wallops.

Support for high I.S. temperatures is provided by Arecibo data (18 N) as shown in Table 7, where biases at $115-120~\rm km$ are about $30-40~\rm K$, that is, about half those at mid-latitudes. The presence of a small component of high energy electrons would give I.S. temperatures in excess of neutral gas temperatures and could possibly account for the discrepancy. The presence of a similar bias in the rocket measurements at Wallops (Table 8), which at 120 km amount to 16 in number (Table 3), means that the discrepancy is unlikely to be solely attributable to biased I.S. temperatures. In particular, the possibility of adjusting to higher N_2 pressures at 130 km should be considered, as by that means higher model temperatures at $115-120~\rm km$ would result and reduce the discrepancy whatever the method of measurement.

Finally, the contrary view to that represented by Case 2 might be taken, namely that the inconsistency is attributable to observed temperatures below 100 km being too high. Case 3 therefore computes temperature models by fitting to data at and above 100 km only and in so doing generates values at 85 - 95 km that are on average ~ 20 K lower than observed values (Table 10). At these heights, temperatures have been measured by the grenade, falling sphere and pitot pressure techniques, which are well-established methods that would not be expected to be consistently biased by more than a few degrees K and certainly not by as much as 20 K. Case 3 is not therefore considered to provide an acceptable model.

No positive recommendation for acceptable models of the intermediate (70 - 130 km) height range has been reached in the above investigations on account of consistent differences between temperature data and the possible models that can be computed with matching conditions at 70 and 130 km. One source of such differences between data and models could be the method of model formulation, and the possibility of incompletely formulated dependences on L.S.T., longitude, and solar activity is pointed out in Section 6. The differences, however, appear to have no consistent relationship with these dependences and are much greater in magnitude than any expected shortcomings in the formulation with respect to these dependences.

Attention therefore needs to be given to the possibility of observed values of temperature being overestimated. Case 2 has been devised to fit observed temperatures closely up to 100 km, but in so doing it provides models above 100 km that are less than observed mid-latitude values by ~80 K at 115 - 120 km (Table 8) in order to match boundary conditions at 70 and 130 km. There may be scope for reducing the ~80 K difference at 115 - 120 km by revising MSIS-86 at 130 km—the region from 130 to, say, 150 km being poorly observed—but the general nature of the inconsistency, that is, observed temperatures in excess of model values, would still remain.

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 <u>Atmospheres, 1979, Environmental Research Papers, No. 677</u>,
 AFGL-TR-79-0241, ADA 081780.

Appendix A

Model Formulation

A1. TEMPERATURE

At height z in the range (z_1, z_2) where $z_1 \le 80$ km (80 km being the upper limit of the lower model^{A1} and $z_2 \ge 90$ km (90 km being the lower limit of the upper model^{A2}), we write

$$T = (M_{Z_1} g/R) W$$
 (A1)

where W is a function of height whose determination is the central consideration of the model formulation. g is acceleration due to gravity, R the universal gas constant and M_{z} the mean molecular weight of air as given by the lower model at height z_1 .

As R may have (slightly) different values for any given lower and upper models and, likewise, g may be represented by different expressions in the two models, the values of R and g at height z are based on a smooth transition from R_1 , $g_1(z)$ for the lower model to R_2 , $g_2(z)$ for the upper model according to the relations

A1. Groves, G.V. (1985) A Global Reference Atmosphere From 18 to 80 km, Air Force Surveys in Geophysics, No. 448, AFGL-TR-85-0129, ADA 162499.

A2. Hedin, A. CIRA 1986 Atmospheric Model in the Region 90 to 2000 km, Draft of 18 June 1986.

$$R = \frac{1}{2} [R_1 + R_2 + (R_2 - R_1) \tanh c\zeta]$$
 (A2)
$$(-1 \le \zeta \le 1)$$

$$g = \frac{1}{2} [g_1 + g_2 + (g_2 - g_1) \tanh c\zeta]$$
 (A3)

where

$$c = 10 \tag{A4}$$

$$\zeta = (z - z_a)/z_d, \quad z_a = \frac{1}{2}(z_1 + z_2), \quad z_d = \frac{1}{2}(z_2 - z_1).$$
 (A5)

With c = 10, the accuracy at $\zeta = \pm 1$ is better than 1 in 10^{-8} . The lower and upper models provide:

$$M_{z_1} = 28.9644 \text{ kg (k mol)}^{-1}, R_1 = 8.31432 \times 10^3, R_2 = 8.314 \times 10^3 \text{ JK}^{-1} \text{ (k mol)}^{-1}$$
(A6)

$$g_1 = g_{\phi}/(1 + z/r_{\phi})^2$$
 at latitude ϕ (A7)

$$g_{\phi} = 9.780356 (1 + 0.0052885 \sin^2 \phi - 0.0000059 \sin^2 2\phi) (ms^{-2})$$
 (A8)

$$r_{\phi} = 2 \times 10^3 \text{ g}_{\phi} / (3.08546 + 0.00227 \cos 2\phi) \text{ (km)}$$
 (A9)

$$g_2 = g_s / (1 + z/R_r)^2$$
 (A10)

where

$$g_s = 9.80665 \text{ m s}^{-2}$$
, $R_r = 6356.77 \text{ km}$. (A11)

(It may be noted that, if the mean molecular weight of air is constant for $z \le z_1$, W is pressure scale height (km) for $z \le z_1$.)

For $z_1 \le z \le z_2$, we express

$$W^{-1} = A + B$$
 (km^{-1}) (A12)

where A and B are polynomials in ζ . B is an interpolating polynomial which depends only on conditions at height z_1 and z_2 , as defined by the particular lower and upper models under consideration, and is otherwise independent of conditions relating to the interval (z_1, z_2) . The conditions imposed at heights z_1 and z_2 are

however such (see below) that B may be considered to provide a first approximation to W^{-1} .

A is an 'adjusting' polynomial whose determination is independent of conditions at heights z_1 and z_2 being dependent on temperature observations in (z_1, z_2) or more correctly on the differences between values of W^{-1} calculated by Eq. (A1) from available temperature values and B calculated for the same geophysical conditions (of date, location, solar activity, and so on).

We take A and its first two height derivatives to be zero at heights \mathbf{z}_1 and \mathbf{z}_2 by expressing it as

$$A = (1 - \zeta^{2})^{3} \sum_{s=1}^{7} \sum_{n=1}^{2} a_{sn} (\zeta^{n} - \gamma_{n}) \xi^{s-1}; \xi = \sin \phi.$$
 (A13)

We choose γ_1 = 0, γ_2 = 1/9 as explained in Appendix A1. The determination of a_{sn} is described in Appendix B, where values of a_{sn} are tabulated.

B is determined by seven conditions involving model values at heights z_1 and z_2 and is therefore taken as a polynomial of degree 6 in ζ

$$B = b_1 + b_2 \zeta + \dots + b_7 \zeta^6 = [1 \zeta \dots \zeta^6] \underline{b}$$
 (A14)

where $\underline{b} = [b_1 \ b_2 \dots b_7]'$. Six of the conditions are for continuity of W^{-1} (and hence of T) and of its first and second height derivatives at heights z_1 and z_2 . The seventh condition is that the ratio of the N_2 pressures at z_1 and z_2 as calculated from the model temperature profile should equal that of the N_2 pressures specified by the lower and upper models at z_1 and z_2 . We write (Appendix A1)

$$\underline{\mathbf{b}} = \underline{\mathbf{S}} \, \underline{\mathbf{\ell}} \qquad (km^{-1}) \tag{A15}$$

where

$$\underline{S} = \begin{bmatrix} 105 & -57 & -57 & -12 & 12 & -1 & 1 \\ 0 & -90 & 90 & -42 & -42 & -6 & 6 \\ -315 & 315 & 315 & 90 & -90 & 9 & 9 \\ 0 & 60 & -60 & 60 & 60 & 12 & -12 \\ 315 & -315 & -315 & -120 & 120 & -15 & -15 \\ 0 & -18 & 18 & -18 & -18 & -6 & 6 \\ -105 & 105 & 105 & 42 & -42 & 7 & 7 \end{bmatrix}$$

$$\underline{\ell} = [\ell_1 \dots \ell_7]' \quad (km^{-1}) \quad . \tag{A17}$$

For the conditions at height z_1 , we have (Appendix A2)

$$\ell_2 = h_1/D_1 + \Delta_1 \tag{A18}$$

$$\ell_4 = (D_2/D_1^2 + \Delta_2)z_d \tag{A19}$$

$$\ell_6 = [(D_1 D_3 + 2 O_2^2)/(h_1 D_1^3) + \Delta_3] z_d^2$$
(A20)

$$D_{1} = 1 + q_{1} \tag{A21}$$

$$D_2 = h_2 - q_2 (A22)$$

$$D_3 = h_1 h_3 - 2 h_2^2 - q_3 \tag{A23}$$

$$h_r = [d^{r-1}(H_{ref}^{-1})/dz^{r-1}]_{z_1}$$
 (r = 1, 2, 3). (A24)

 H_{ref} is the zonal mean pressure scale height of the lower model and hence from Eq. (A24) and Reference A1 we have

$$h_1 = \sum_{n=1}^{9} \sum_{s=1}^{9} c_{ns} \xi^{s-1} Z_1^{n-1}, \qquad \xi = \sin \phi \ (km^{-1})$$
 (A25)

$$h_2 = \sum_{n=1}^{9} \sum_{s=1}^{9} (n-1) c_{ns} \xi^{s-1} Z_1^{n-2} / Z_d \qquad (km^{-2})$$
 (A26)

$$h_3 = \sum_{n=1}^{9} \sum_{s=1}^{9} (n-1) (n-2) c_{ns} \xi^{s-1} Z_1^{n-3} / Z_d^2 (km^{-3})$$
(A27)

where

$$Z_1 = (z_1 - Z_2)/Z_d$$
, $Z_2 = 48.75 \text{ km}$, $Z_d = 31.25 \text{ km}$ (A28)

and c_{ns} are linearly interpolated to the required date from the values tabulated in units of km⁻¹ in Reference 1. The dependence of l_2 , l_4 , l_6 on longitude, λ , is expressed by q_r , which by Eq. (A2.3) and Eq. (A2.7) is

$$q_r = K_{1r} \cos \lambda + L_{1r} \sin \lambda + K_{2r} \cos 2\lambda + L_{2r} \sin 2\lambda$$
 (A29)

where, for j = 1, 2; r = 1, 2, 3,

$$K_{jr} = h_1^r Y_r, \qquad Y_r = (d^{r-1}y/dz^{r-1})_{z_1}$$
(A30)

$$y = R_1 T_j \frac{\cos \lambda_{T_j}}{\sin \lambda_{T_j}} \lambda_{T_j} (km)$$
 (A31)

Values of y are calculated from the values of T_j , λ_{Tj} tabulated in Reference 1 for each month at 10° latitude steps and are linearly interpolated to the required date and latitude. Y_r are obtained numerically from values y_{-1/2}, y_{1/2}, y_{3/2}, y_{5/2} of y at four equally-spaced heights $z_{-1/2}$, $z_{1/2}$, $z_{3/2}$, $z_{5/2}$ of which z_1 is the mid-point and Δz is the increment step by fitting a cubic to provide

$$Y_1 = (q a_+ - b_+)/16$$
 (A32)

$$Y_2 = (27 a_- b_-)/24 \Delta z$$
 (A33)

$$Y_3 = (b_+ - a_+)/2 (\Delta z)^2$$
 (A34)

where

$$a_{\pm} = y_{3/2} \pm y_{1/2}$$
 $b_{\pm} = y_{5/2} \pm y_{-1/2}$ (A35)

For the tabulations in Reference A1, Δ_z = 4 km. For z_1 = 70 km (the value adopted), the four equally-spaced heights are 64, 68, 72, and 76 km. (For other values of z_1 , Eq. (A32) to Eq. (A34) may need to be replaced by alternative formulas.) The dimensions of q_r are (km) $^{-2(r-1)}$ for r=1,2,3.

The terms Δ_r relate to incremental changes of temperature in the vicinity of height z_1 that may be associated with the solar cycle. Corresponding to a change ΔR_n in sunspot number from a reference value R_{no} , we have the formulation (Appendix A2)

A3. Groves, G.V. (1986) An empirical model for solar cycle changes in mesospheric structure at longitudes 44 - 77°E, Planet. Space Sci. 34:1037-1041.

$$\Delta_{r} = -K(\phi) \Phi_{r} \Delta R_{n} \quad (r = 1, 2, 3) \tag{A36}$$

where

$$K(\phi) = p + q \cos^{n} \phi \tag{A37}$$

and

$$\Phi_1 = (1 - \tau^2) (1 + \alpha \tau^3) \tag{A38}$$

$$\Phi_2 = \tau (1 - \tau^2) (-2 + 3 \alpha \tau - 5 \alpha \tau^3) / \ell \tag{A39}$$

$$\Phi_3 = 2(1-\tau^2) \left[(1-3\tau^2) + \alpha \tau (3-16\tau^2+15\tau^4) \right] / \ell^2$$
 (A40)

where

$$\tau = \tanh[(z_1 - a)/\ell] . \tag{A41}$$

The quantities α , a, ℓ , p, q and n are calculated from

$$\theta = \theta_1 + f \theta_2$$
 $(\theta = \alpha, a, \ell, p, q, or n)$ (A42)

$$f = \tanh(4\phi/\pi)\cos[\pi(t_d - 1)/182.5]$$
 (A43)

 t_d being the day number in the year. Numerical values for θ_1 , θ_2 ($\theta = \alpha$, a, ℓ , p, q or n) have been given but are tentative, being derived from limited data at eastern longitudes and therefore this dependence has only been included in exploratory calculations that are not a part of this report. The model of Reference A1 is for data over years of average sunspot number, $R_{n0} = 65$.

For the conditions at height z_2 , we have (Appendix A3)

$$\ell_3 = 10^3 \text{ M}_{z_1} \text{ g}_2 (z_2)/\text{R}_2 \text{ T}(z_2)$$
 (A44)

where $T(z_2)$ is the temperature at height z_2 calculated from MSIS-86 as defined in Reference A2 for the required time, date, latitude, longitude, solar activity and so on. Also

$$\ell_5 = -(2d + u) \ell_3 z_d$$
 (A45)

$$\ell_7 = (6 d^2 + 6 ud + \eta) \ell_3 z_d^2$$
 (A46)

where

$$d = (R_p + z_2)^{-1}$$
 (km⁻¹). (A47)

For $z_2 \ge z_a = 117.2 \text{ km}$

$$u = -\sigma_g [1 - T_\infty / T(z_2)]$$
 (km⁻¹) (A48)

$$\eta = u(2u + \sigma_g) \qquad (km^{-2})$$
 (A49)

$$\sigma_g = \sigma[g_2(z_2)/g_2(z_\ell)]$$
 (km⁻¹) (A50)

where σ is taken from MSIS-86, being related to the MSIS-86 parameters of temperature T_{ℓ} and temperature gradient T'_{ℓ} at z_{ℓ} = 120 km and the MSIS-86 exospheric temperature T_{∞} by $\sigma = T'_{\ell}/(T_{\infty} - T_{\ell})$ as given in Reference A2.

For $z_2 \leq z_a = 117.2 \text{ km}$,

$$u = -2\chi_2 T(z_2)(T_B + 2T_C \chi_2^2 + 3T_D \chi_2^4) (d\chi/dz)_{z_2}$$
(A51)

$$\eta = 2 \text{ T}(z_2) (T_B + 6 T_C \chi_2^2 + 15 T_D \chi_2^4) (d\chi/dz)_{z_2}^2$$
 (A52)

where

$$\chi_2 = -[\xi (z_2, z_a) - \xi (z_0, z_a)]/\xi (z_0, z_a)$$
(A53)

with

$$\xi(z, z_a) = (z - z_a) (R_r + z_a) / (R_r + z)$$
 (A54)

and

The coefficients T_B , T_C , T_D are taken from MSIS-86 being related to the MSIS-86 parameters of temperature T_a and temperature gradient T_a' at $z=z_a$ and the MSIS-86 temperature T_o at height z_o of the MSIS-86 mesopause.

Finally, for the seventh condition, which involves values at both heights z_1 and z_2 , we have (Appendix A4)

$$\ell_1 = (M_{z_1}/M_h) X/z_d$$
 (km⁻¹) (A56)

where $M_h = 28 \text{ kg/(kmol)}^{-1}$ and X is obtained by iteration (over only two or three cycles for 10^{-10} accuracy on taking an initial value of $X = (z_2 - z_1)/(7 \text{ km})$ from Eq. (A4.9) written as

$$X = (M_h/\overline{M}_0) \left[\ln \mu_{N_2} - A_{N_2}^{-1} \ln (1 + \nu_{N_2}^{A_{N_2}}) - e^{-X} \right]$$
 (A57)

where $\overline{M}_0 = 28.95 \text{ kg(kmol)}^{-1}$ and

$$\mu_{N_2} = f_{N_2} p_1(z)/k(z_2) p_m(z_2, M_{N_2});$$
 $k(z_2) = 1.000292$ (A58)

$$\nu_{N_2} = n_d (z_2, M_{N_2}) / n_m (z_2, M_{N_2})$$
 (A59)

$$A_{N_2} = M_h / (\overline{M}_o - M_{N_2})$$
 (A60)

where $M_{N_2} = 28 \text{ kg(kmol)}^{-1}$. f_{N_2} is the N_2 fractional volume of air at height z_1 (= 0.78084) and $p(z_1)$, the air pressure at height z_1 , is

$$p(z_1) = p_{ref}(z_1) (1 + D_0) (1 + \Delta_0)$$
(A61)

where $p_{ref}(z_1)$ is the zonal mean air pressure at height z_1 calculated from Reference A1

$$p_{ref}(z_1) = exp(-31.25 \sum_{n=0}^{9} \sum_{s=1}^{9} c_{ns} \xi^{s-1} \zeta^n/_n)$$
 (mb) (A62)

 $(\zeta^n)_n$ denoting unity for n=0). c_{ns} are linearly interpolated to the required date from the values tabulated in Reference A1. Dependence of $p(z_1)$ on longitude λ is introduced through

$$D_0 = K_{10} \cos \lambda + L_{10} \sin \lambda + K_{20} \cos 2\lambda + L_{20} \sin 2\lambda$$
 (A63)

where, for j = 1, 2,

$$K_{j0} = (y)_{z_1}$$
, $y = p_j \cos \lambda_{pj}$. (A64)

Values of y are calculated from the values of p_j , λ_{pj} tabulated in Reference A1 for each month and 10° latitude step and are linearly interpolated to the required date and latitude. The value of y at height z_1 is obtained by an interpolating cubic through four values $y_{-1/2}$, $y_{1/2}$, $y_{3/2}$, $y_{5/2}$ at heights $z_{-1/2}$, $z_{3/2}$, $z_{5/2}$ (as for y_1 above), that is,

$$(y)_{z_1} = (9 a_+ - b_+)/16$$
 (A65)

where

$$a_{+} = y_{3/2} + y_{1/2}$$
 $b_{+} = y_{5/2} - y_{-1/2}$ (A66)

(For values of z_1 other than 70 km (the adopted value), Eq. (A66) may need to be replaced by an alternative formula.) Δ_0 is a relative pressure increase at height z_1 that may be associated with the solar cycle corresponding to Eq. (A36). We have

$$\Delta_0 = K(\phi) \ell \Phi_0 \Delta R_n \tag{A67}$$

where

$$\Phi_0 = 1 + \tau + \frac{\alpha}{4} (\tau^4 - 1) . \tag{A68}$$

 n_d (z_2 , M_{N_2}), the N_2 diffusive profile number density, and n_m (z_2 , M_{N_2}), the N_2 mixing profile number density, at height z_2 that appear in Eq. (A59) are MSIS-86 parameters that are calculated in units of cm⁻³ by the MSIS-86 formulation of Reference A2 for the required time, date, latitude, longitude, solar activity, and so on. In units of mb, we have

$$p_{m}(z_{2}, M_{N_{2}}) = 10^{4} R_{2} A_{N_{2}}^{-1} n_{m}(z_{2}, M_{N_{2}}) T(z_{2})$$
(A69)

where $T(z_2)$ is the MSIS-86 temperature at height z_2 for the required conditions and A_{N2}^{-1} is the reciprocal of Avogadro's number used in MSIS-86, that is,

$$A_{N2}^{-1} = 1.66 \times 10^{-27}$$
 (kmol) . (A70)

Hence $A_{N2} = 6.02410 \times 10^{26} \text{ (kmol)}^{-1}$.

A2. DENSITY

At height z in the range (z₁, z₂) density ρ is calculated from

$$\rho = \sum_{i \neq 6} M_i n(z, M_i) / A_N$$
 (kg m⁻³) (A71)

where $n(z, M_1)$ is the number density (m^{-3}) at height z of gas constituent of molecular weight M_1 , the values of M_1 being taken to be those of MSIS-86, that is, $M_1 = 4$, $M_2 = 16$, $M_3 = 28$, $M_4 = 32$, $M_5 = 40$, $M_7 = 1$ and $M_8 = 14$ corresponding to gases He, O, N_2 , O_2 , Ar, H and N. When different values of Avogadro's number are adopted for the lower and upper models, we take a smooth transition given by

$$A_{N} = \frac{1}{2} [A_{N1} + A_{N2} + (A_{N2} - A_{N1}) \tanh c\zeta] \qquad (-1 \le \zeta \le 1)$$
 (A72)

where

$$A_{N1} = 6.02257 \times 10^{26}$$
 $A_{N2} = 6.02410 \times 10^{26}$ (mks units) (A73)

in the present calculation. For $n(z, M_i)$ we adopt the relations of the MSIS-86 formulation of Reference A2 and have (Appendix A5)

$$n(z_1, M_i) = 10^6 [n_d(z, M_i)^{A_i} + n_m(z, M_i)^{A_i}]^{A_i} C_{1i}(z) C_{2i}(z) (m^{-3})$$
 (A74)

$$A_{i} = M_{h}/(\overline{M}_{o} - M_{i}) \tag{A75}$$

$$n_{d}(z, M_{i}) = n_{d}(z_{2}, M_{i}) e^{M_{i}J(\zeta)} [T(z_{2})/T(z)]$$
 (A76)

$$n_{m}(z, M_{i}) = n_{m}(z_{2}, M_{i}) e^{M_{o}J(\zeta)}$$
 T(z₂)/T(z) (A77)

$$J(\zeta) = [U(1) - U(\zeta)] z_d / M_{z_1}$$
 (A78)

$$U(\zeta) = \sum_{s=1}^{7} \sum_{n=1}^{2} a_{sn} \phi_{n}(\zeta) \xi^{s-1} + b_{1} \zeta + \frac{1}{2} b_{2} \zeta^{2} + \dots + \frac{1}{7} b_{7} \zeta^{7}$$
(A79)

$$\phi_{n}(\zeta) = \zeta [\psi_{1}(\zeta, n) - \gamma_{n} \psi_{2}(\zeta)]$$
 (A80)

$$\psi_1(\zeta, n) = \zeta^n \left(\frac{1}{n+1} - \frac{3\zeta^2}{n+3} + \frac{3\zeta^4}{n+5} - \frac{\zeta^6}{n+7} \right)$$
 (A81)

$$\psi_2(\zeta) = 1 - \zeta^2 + \frac{3}{5} \zeta^4 - \frac{1}{7} \zeta^6 \tag{A82}$$

where α_i = -0.4 for i = 1 and 7 (He and H) and is otherwise zero. C_{1i} = 1 for N_2 (i = 3) and C_{2i} = 1 for He, N_2 , O_2 and Ar (i = 1, 3, 4 and 5). For the remaining cases

$$C_{ji}(z) = \exp \left\{ r_{ji} / [1 + \exp(z - z_{ji}) / H_{ji}] \right\}$$
 (j = 1, 2) (A83)

$$r_{1i} = \frac{1}{2} [(R'_{1i} + R_{1i}) + (R_{1i} - R'_{1i}) \tanh c\zeta]$$
 (A84)

$$\mathbf{r_{2i}} = \mathbf{R_{2i}} \tag{A85}$$

where

$$R_{1i} = \ln \left[r_i \, m_m (z_\ell, M_{N_2}) / n_m (z_\ell, M_i) \right] .$$
 (A86)

The values of r_i , z_{ji} , H_{ji} , R_{2i} are those of MSIS-86 (Reference A2, Table 2d) and for i = 2, 7 and 8 (O, H and N)

$$R'_{1i} = R_{1i} \tag{A87}$$

whereas for i = 1, 4 and 5 (He, O_2 and Ar), $R_{1i}^{'}$ is chosen so that these constituents have been given volume fractions of air, f_i , at height z_1 . The required values are given by (Appendix A5)

$$R'_{1i} = \left[\ln \mu_{i} - A_{i}^{-1} \ln \left\{ \left[\nu_{i} \stackrel{/}{e}^{M_{i} J(-1)} \left(T(z_{2}) / T(z_{1}) \right)^{\alpha_{i}} \right]^{A_{i}} \left[e^{\overline{M}_{0}(J-1)} \right]^{A_{i}} \right\} \right]$$

$$\left\{ 1 + \exp \left[(z_{1} - z_{1i}) / H_{1i} \right] \right\}$$
(A88)

$$p_{m}(z_{2}, M_{i}) = R_{2} A_{N2}^{-1} n_{m}(z_{2}, M_{i}) T(z_{2})$$
 (A89)

$$\mu_{i} = f_{i} p(z_{1})/\kappa (z_{2}) p_{m}(z_{2}, M_{N_{2}})$$
 (A90)

$$v_i = n_d (z_2, M_i) / n_m (z_2, M_i)$$
 (A91)

for i = 1, 4 and 5, where we take

$$f_1 (\equiv f_{He}) = 5.24 \times 10^{-6}, \quad f_4 (\equiv f_{O_2}) = 0.21023, \quad f_s (\equiv f_{Ar}) = 9.34 \times 10^{-3}$$

$$\kappa(z_2) = \frac{8.31432 \times 6.02410}{6.02257 \times 8.314} = 1.000292 . \quad (A92)$$

In order to maintain continuity in the mean molecular weight at height z_1 , the value taken for f_{0_2} is such that the mean molecular weight of air at height z_1 for the gas constituents N_2 , O_2 , Ar, He is equal to M_{z_1} , that is, such that $28 \, f_{N_2} + 32 \, f_{0_2} + 40 \, f_{Ar} + 4 \, f_{He} = M_{z_1}$ (= 28.9644). The value obtained, 0.21023, is slightly higher than the actual value, 0.20948, the excess being largely accounted for by the presence of CO_2 in the real atmosphere which is omitted in the model.

A useful check on computing accuracy is obtained by evaluating $R_{1i}^{'}$ for i=3 (that is, N_2) and comparing its value with the adopted value of zero.

A3. PRESSURE

Total number density is obtained as

$$n = \sum_{m \neq 6} n(z, M_i)$$
 (m⁻³). (A93)

Mean molecular weight is then

$$\overline{M} = A_N \rho/n \qquad (kg (k mol)^{-1}) \qquad (A94)$$

and pressure is

$$p = (R/A_N) n T$$
 (N m⁻²). (A95)

A4. TURBOPAUSE HEIGHT \mathbf{z}_{hi} OF THE i th GAS CONSTITUENT

In the MSIS-86 formulation, $^{\mathrm{A2}}$ turbopause height is taken to be the height z_{hi} at which

$$n_{m}(z_{hi}, M_{i}) = n_{d}(z_{hi}, M_{i})$$
(A96)

and z_{hi} is a fixed parameter, being 105 km for N_2 , O_2 , Ar, N and O, 100 km for He and 95 km for H. In the present formulation, where MSIS-86 parameters are adopted for the height z_2 which (at 130 km) exceeds the heights z_{hi} , the MSIS-86 values for z_{hi} cannot hold unless the derived temperature profile in (z_{hi}, z_2) is the same as the MSIS-86 profile, which clearly is not the case. However, as deviations between the two profiles are not large, the values of z_{hi} satisfying Eq. (A96) would not be expected to differ greatly from the MSIS-86 values. In the computations that have been undertaken the differences have been less than 1 km.

We calculate

$$z_{hi} = z_a + z_d \zeta_{hi}$$
 (A97)

where

$$\zeta_{hi} = \left\{ U(1) - \frac{z_d^{-1} M_{z_1}}{(M_o - M_i)} \ell_n \left[\nu_i \left(\frac{T(z_2)}{T(z_1)} \right)^{\alpha_i} \right] \right\} / \left[U(\zeta_{hi})/\zeta_{hi} \right]. \quad (A98)$$

The denominator in Eq. (A98) is calculated from

$$[U(\zeta)/\zeta] = \sum_{s=1}^{7} \sum_{n=1}^{2} a_{sn} \left[\psi_{1}(\zeta_{i}, n) - \gamma_{n} \psi_{2}(\zeta) \right] \xi^{s-1} + b_{1} + \frac{1}{2} b_{2} \zeta + \dots + \frac{1}{7} b_{7} \zeta^{6} . \tag{A99}$$

 ζ_{hi} is obtained from Eq. (A98) by iteration with a suitable initial value (corresponding to say z_{hi} = 100 km).

Appendix A1

Formulation of W⁻¹

We define

$$W^{-1} = A + B$$
 $(-1 \le \xi \le 1; -1 \le \zeta \le 1)$ (A1.1)

where, for given S, N and a $_{\mbox{sn}}$ (which are independent of ξ and $\zeta)$

$$A = (1 - \zeta^{2})^{3} \sum_{s=1}^{S} \sum_{n=1}^{N} a_{sn} (\zeta^{n} - \gamma_{n}) \xi^{s-1}$$
(A1.2)

$$B = \sum_{r=1}^{7} b_r \zeta^{r-1} . (A1.3)$$

 $\gamma_{\,\,n}$ are chosen so that for all values of ξ

$$\int_{-1}^{1} A \, d\zeta = 0 \quad . \tag{A1.4}$$

Hence

$$\gamma_{n} = G_{n}/G_{o} \tag{A1.5}$$

where

$$G_{n} = \int_{-1}^{1} (1 - \zeta^{2})^{3} \zeta^{n} d\zeta.$$
 (A1.6)

On integrating by parts it may be shown that

$$(n+7) G_n = (n-1) G_{n-2}$$
 (A1.7)

Hence for n even

$$\gamma_n = (G_n/G_{n-2}) \dots (G_2/G_0)$$

$$= \frac{1 \cdot 3 \cdot \dots (n-3) (n-1)}{9 \cdot 11 \cdot \dots (n+5) (n+11)} \quad (n \text{ even}) \quad . \tag{A1.8}$$

For n odd, $G_1 = 0$ by direct evaluation and hence $G_3 = G_5 = \dots = 0$ and

$$\gamma_n = 0 \quad (n \text{ odd}) \quad .$$
 (A1.9)

 $\mathbf{b_r}$ are chosen so that for given ℓ_1 , ..., ℓ_7 , by Eq. (A1.1) to Eq. (A1.4),

$$\int_{-1}^{1} B d\zeta = \int_{-1}^{1} W^{-1} d\zeta = \ell_{1}$$
 (A1.10)

$$B(-1) = W^{-1}(-1) = \ell_2, \quad \frac{dB(-1)}{d\zeta} = \frac{dW}{d\zeta}^{-1}(-1) = \ell_4, \quad \frac{d^2B(-1)}{d\zeta^2} = \frac{d^2W}{d\zeta^2}^{-1}(-1) = \ell_6$$
(A1.11)

$$B(1) = W^{-1}(1) = \ell_3, \quad \frac{dB(1)}{d\zeta} = \frac{dW}{d\zeta}^{-1}(1) = \ell_5, \quad \frac{d^2B(1)}{d\zeta^2} = \frac{d^2W}{d\zeta^2}^{-1}(1) = \ell_7. \tag{A1.12}$$

These conditions require that

$$\underline{S}^{-1}\underline{b} = \underline{\ell} \tag{A1.13}$$

where
$$\underline{\mathbf{b}} = \begin{bmatrix} \mathbf{b}_1 & \cdots & \mathbf{b}_7 \end{bmatrix}'$$
, $\underline{\boldsymbol{\ell}} = \begin{bmatrix} \boldsymbol{\ell}_1 & \cdots & \boldsymbol{\ell}_7 \end{bmatrix}'$ and

$$\underline{S}^{-1} = \begin{bmatrix} 2 & 0 & \frac{2}{3} & 0 & \frac{2}{5} & 0 & \frac{2}{7} \\ 1 & -1 & 1 & -1 & 1 & -1 & 1 \\ 1 & 1 & 1 & 1 & 1 & 1 & 1 \\ 0 & 1 & -2 & 3 & -4 & 5 & -6 \\ 0 & 1 & 2 & 3 & 4 & 5 & 6 \\ 0 & 0 & 2 & -6 & 12 & -20 & 30 \\ 0 & 0 & 2 & 6 & 12 & 20 & 30 \end{bmatrix}$$
(A1.14)

From Eq. (A1.13), we obtain

$$\underline{\mathbf{b}} = \underline{\mathbf{S}} \, \underline{\boldsymbol{\ell}} \tag{A1.15}$$

where, on inverting \underline{S}^{-1}

$$\underline{S} = \begin{bmatrix} 105 & -57 & -57 & -12 & 12 & -1 & 1 \\ 0 & -90 & 90 & -42 & -42 & -6 & 6 \\ -315 & 315 & 315 & 90 & -90 & 9 & 9 \\ 0 & 60 & -60 & 60 & 60 & 12 & -12 \\ 315 & -315 & -315 & -120 & 120 & -15 & -15 \\ 0 & -18 & 18 & -18 & -18 & -6 & 6 \\ -105 & 105 & 105 & 42 & -42 & 7 & 7 \end{bmatrix}$$

Appendix A 2

Conditions at Height z₁

By definition

$$W^{-1} = \frac{M_{z_1} g}{R T} = \frac{M_{z_1} g}{R T_{ref}} \frac{T_{ref}}{T(\lambda)}$$
 (A2.1)

where the dependence of T on longitude λ is now indicated and T_{ref} denotes the zonal mean of $T(\lambda)$.

At height z_1 , which is at or below 80 km, we have $R = R_1$, $g = g_1$ (as defined by Eqs. (A6) and (A7)) and from Reference B1, Appendix B.

$$H_{ref}^{-1} = M_{z_1} g_1 / R_1 T_{ref} = \sum_{n=1}^{9} \sum_{s=1}^{9} c_{ns} \xi^{s-1} Z^{n-1}$$
 (A2.2)

$$\Delta T = T(\lambda) - T_{ref} = T_1 \cos(\lambda - \lambda_{T1}) + T_2 \cos(2\lambda - \lambda_{T2})$$
 (A2.3)

where ξ = sin (latitude), $Z = (Z - Z_a)/Z_d$, $Z_a = 48.75$ km, $Z_d = 31.25$ km and c_{ns} , T_1 , T_2 , λ_{T1} , λ_{T2} are tabulated in Reference B1. Hence from Eq. (A2.1) to Eq. (A2.3)

$$W^{-1} = h_1/D_1$$
 (A2.4)

$$dq_2/dz = (q_3 + 2h_2 q_2)/k_1$$
 (A2.14)

Hence Eq. (A2.12) to Eq. (A2.14) give, on using Eq. (A2.11)

$$d^{2}W^{-1}/dz^{2} = (D_{1}D_{3} + 2D_{2}^{2})/(h_{1}D_{1}^{3})$$
(A2.15)

where

$$D_3 = h_1 h_3 - 2 h_2^2 - q_3 . (A2.16)$$

The required conditions at height z_1 , which are expressed by Eq. (A1.11), become by Eqs. (A2.4), (A2.10), (A2.15) and (A5)

$$\ell_2 = h_1/D_1 + \Delta_1$$
 (A2.17)

$$\ell_4 = (D_2/D_1^2 + \Delta_2) z_d$$
 (A2.18)

$$\ell_6 = \left[(D_1 D_3 + 2 D_2^2) / (h_1 D_1^3) + \Delta_3 \right] z_d^2$$
 (A2.19)

where h_r , q_r and D_r are evaluated at $z = z_1$ and the terms

$$\Delta_{r} = \left[d^{r-1} (\Delta W^{-1})/dz^{r-1} \right]_{z_{1}}$$
 (A2.20)

are introduced to enable an imposed change ΔW^{-1} of W^{-1} in the vicinity of $z=z_1$ to be taken into account due, for example, to a dependence on solar activity. Corresponding to an increase in temperature ΔT

$$\Delta \mathbf{W}^{-1} = - \mathbf{W}^{-1} \mathbf{T}^{-1} \Delta \mathbf{T} . \tag{A2.21}$$

A formulation has been given in Reference A3 for Δ T corresponding to a change in sunspot number Δ R_n, by which Eqs. (A2.20) and (A2.21) yield

$$dq_2/dz = (q_3 + 2h_2 q_2)/k_1$$
 (A2.14)

Hence Eq. (A2. 12) to Eq. (A2. 14) give, on using Eq. (A2. 11)

$$d^{2}W^{-1}/dz^{2} = (D_{1}D_{3} + 2D_{2}^{2})/(h_{1}D_{1}^{3})$$
(A2.15)

where

$$D_3 = h_1 h_3 - 2 h_2^2 - q_3 . (A2.16)$$

The required conditions at height z_1 , which are expressed by Eq. (A1.11), become by Eqs. (A2.4), (A2.10), (A2.15) and (A5)

$$\ell_2 = h_1/D_1 + \Delta_1$$
 (A2.17)

$$\ell_4 = (D_2/D_1^2 + \Delta_2) z_d$$
 (A2.18)

$$\ell_6 = \left[(D_1 D_3 + 2 D_2^2) / (h_1 D_1^3) + \Delta_3 \right] z_d^2$$
 (A2.19)

where h_{r} , q_{r} and D_{r} are evaluated at $z = z_{1}$ and the terms

$$\Delta_{r} = \left[d^{r-1} (\Delta W^{-1})/dz^{r-1} \right]_{z_{1}}$$
 (A2.20)

are introduced to enable an imposed change ΔW^{-1} of W^{-1} in the vicinity of $z=z_1$ to be taken into account due, for example, to a dependence on solar activity. Corresponding to an increase in temperature ΔT

$$\Delta \mathbf{W}^{-1} = - \mathbf{W}^{-1} \mathbf{T}^{-1} \Delta \mathbf{T} . \tag{A2.21}$$

A formulation has been given in Reference A3 for Δ T corresponding to a change in sunspot number $\Delta\,R_n$, by which Eqs. (A2. 20) and (A2. 21) yield

$$\Delta_{\mathbf{r}} = -K(\phi) \Phi_{\mathbf{r}} \Delta R_{\mathbf{n}}$$
 (A2.22)

$$K(\phi) = p + q \cos^{n} \phi \tag{A2.23}$$

$$\Phi_{\mathbf{r}} = \frac{d^{\mathbf{r}-1}}{dz^{\mathbf{r}-1}} \left[\frac{1 + \alpha \tanh^3 Z_{\mathbf{s}}}{\cosh^2 Z_{\mathbf{s}}} \right]_{z=z_1}$$
(A2.24)

where

$$Z_{s} = (z-a)/\ell$$
 . (A2.25)

The quantities α , a, ℓ , p, q and n are calculated from

$$\theta = \theta_1 + f \theta_2$$
 ($\theta = \alpha$, a, ℓ , p, q or n) (A2.26)

$$f = \tanh (4\phi/\pi) \cos [\pi(t_d - 1)/182.5]$$
 (A2.27)

where t_d is the day number in the year and $\alpha_1 = 0.60$, $a_1 = 66.63$ km, $\ell_1 = 12.14$ km, $p_1 = 7.19 \times 10^{-5}$ km⁻¹, $q_1 = 4.80 \times 10^{-5}$ km⁻¹, $n_1 = 6.0$, $\alpha_2 = 0$, $a_2 = -3.09$ km, $\ell_2 = 0$, $p_2 = -0.51 \times 10^{-5}$ km⁻¹, $q_2 = -2.16 \times 10^{-5}$ km⁻¹, $n_2 = -5.0$.

Appendix A 3

Conditions at Height z

At height $z_2 (\ge 90 \text{ km})$, MSIS-86 is valid and we have $R = R_2$, $g = g_2$ (as defined by Eq. (A.6) and Eq. (A.10)) and Eq. (A.1) becomes

$$W^{-1} = M_{z_1} g_2/R_2 T$$
 (A3.1)

By logarithmic differentiation

$$(W^{-1})^{-1} d(W^{-1})/dz = -2d - T^{-1} dT/dz$$
 (A3.2)

where

$$d = 1/(R_p + z)$$
 (A3.3)

A further differentiation gives, by Eq. (A3.3),

$$(W^{-1})^{-1} \frac{d^2 W^{-1}}{dz^2} - (W^{-1})^{-2} \left(\frac{dW}{dz}^{-1} \right)^2 = 2 d^2 + \left(\frac{1}{T} \frac{dT}{dz} \right)^2 - \frac{1}{T} \frac{d^2 T}{dz^2} . \quad (A3.4)$$

Hence from Eqs. (A3.2) and (A3.4)

$$(W^{-1})^{-1} \frac{d^2 W^{-1}}{dz^2} = 6 d^2 + 4 d \frac{1}{T} \frac{dT}{dz} + 2 \left(\frac{1}{T} \frac{dT}{dz} \right)^2 - \frac{1}{T} \frac{d^2 T}{dz^2} .$$
 (A3.5)

From Reference A2, different expressions hold for dT/dz, d^2T/dz^2 according to whether z is greater or less than $z_a = 117.2$ km.

For $z_2 \stackrel{?}{\geq} z_a$, we have A^2

$$T(z) = T_{\infty} - (T_{\infty} - T_{\ell}) \exp[-\sigma \xi (z, z_{\ell})]$$
 (A3.6)

$$\xi(z, z_{\ell}) = (z - z_{\ell}) (R_p + z_{\ell})/(R_p + z)$$
 (A3.7)

where T_{∞} is exospheric temperature, T_{ℓ} is temperature at height z_{ℓ} (= 120 km) and R_{p} = 6356.77 km. σ is a constant that is related to the temperature gradient, T_{ℓ} , at $z = z_{\ell}$ by $\sigma = T_{\ell}^{\prime}/(T_{\infty} - T_{\ell})$. From Eq. (A3.6), we have

$$\frac{1}{T} \frac{dT}{dz} = -\sigma_g \left(1 - \frac{T_{\infty}}{T} \right) \tag{A3.8}$$

$$\frac{1}{T}\frac{d^2T}{dz^2} = -\sigma_g\left(1 - \frac{T_\infty}{T}\right)\left[\frac{d}{dz}\left(\ln\frac{d\xi}{dz}\right) - \sigma_g\right] \tag{A3.9}$$

where

$$\sigma_{g} = \sigma \, d\xi / dz \tag{A3.10}$$

and from Eq. (A3.7), by Eqs. (A10) and (A3.3),

$$d\xi /dz = (R_p + z_{\ell})^2 / (R_p + z)^2 = g_2(z) / G_2(z_{\ell})$$
(A3.11)

$$\frac{\mathrm{d}}{\mathrm{d}z} \left[\ln \left(\frac{\mathrm{d}\xi}{\mathrm{d}z} \right) \right] = -2\mathrm{d} \quad . \tag{A3.12}$$

Hence if

$$u = -\sigma_{g} (1 - T_{\infty}/T)$$
 (A3.13)

$$\eta = u (2 u + \sigma_g).$$
(A3.14)

Equations (A3.2) and (A3.5) give by Eqs. (A3.8) to (A3.13)

$$(W^{-1})^{-1} dW^{-1}/dz = -(2d + u)$$
 (A3.15)

$$(W^{-1})^{-1} d^2W^{-1}/dz^2 = 6 d^2 + 6 u d + \eta$$
 (A3.16)

where by Eqs. (A3.10) and (A3.11)

$$\sigma_{g} = \sigma g_{2}(z)/g_{2}(z_{\ell})$$
 (A3.17)

For $z_2 \le z_a$, we have from Reference A2

$$T^{-1} = T_{O}^{-1} + T_{B}x^{2} + T_{C}x^{4} + T_{D}x^{6}$$
(A3.18)

where

$$\mathbf{x} = -\left[\xi (\mathbf{z}, \mathbf{z}_{a}) - \xi (\mathbf{z}_{o}, \mathbf{z}_{a})\right] / \xi (\mathbf{z}_{o}, \mathbf{z}_{a}) \tag{A3.19}$$

$$\xi(z, z_a) = (z - z_a) (R_p + z_a)/(R_p + z)$$
 (A3.20)

 z_0 is mesopause height as determined by the MSIS-86 formulation. By Eq. (A3.18) we have

$$\frac{1}{T} \frac{dT}{dz} = -T \frac{dT^{-1}}{dz} = -2xT (T_B + 2T_C x^2 + 3T_D x^4) \frac{dx}{dz}$$
 (A3.21)

$$2\left(\frac{1}{T} \frac{dT}{dz}\right)^{2} - \frac{1}{T} \frac{d^{2}T}{dz^{2}} = T \frac{d^{2}T}{dz^{2}}^{-1} = 2 \times T \left(T_{B} + 2 T_{C} \times^{2} + 3 T_{D} \times^{4}\right) \frac{dx}{dz} \frac{d}{dz} \left[\ln \left(\frac{dx}{dz}\right) \right]$$

+ 2 T (T_B + 6 T_C
$$x^2$$
 + 15 T_D x^4) $\left(\frac{dx}{dz}\right)^2$. (A3.22)

Equations (A3.2) and (A3.5) can then be written, by Eq. (A3.19) to Eq. (A3.22) and Eq. (A3.3), as Eq. (A3.15) and (A3.16), where

$$u = -2 \times T (T_B + 2 T_C x^2 + 3 T_D x^4) \frac{dx}{dz}$$
 (A3.23)

$$\eta = 2 \text{ T}(T_B + 6 T_C x^2 + 15 T_D x^4) \left(\frac{dx}{dz}\right)^2$$
 (A3.24)

and by Eqs. (A3.19), (A3.20) and (A10)

$$\frac{dx}{dz} = -\frac{1}{\xi (z_0, z_a)} \frac{g_2(z)}{g_2(z_a)}.$$
 (A3.25)

The required conditions at height z_2 , which are expressed by Eq. (A1.12), become by Eqs. (A3.1), (A3.15) and (A3.16), on changing units from m to km in ℓ_3

$$\ell_3 = 10^3 \text{ M}_{z_1} g_2(z_2)/R_2 T(z_2)$$
 (km⁻¹) (A3.26)

$$\ell_5 = -(2d + u) \ell_3 z_d$$
 (km⁻¹) (A3.27)

$$l_7 = (6 d^2 + 6 ud + \eta) l_z z_d^2$$
 (km⁻¹) (A3.28)

where d, u and η are evaluated at z = z_2 .

Appendix A4 The N Pressure Ratio Condition

We adopt the MSIS-86 representation of Reference A2 for the number density profile of a gas constituent, n(z, M) as a meld of a diffusive profile $n_d(z, M)$ and a mixing profile $n_m(z, M)$. For molecular nitrogen, $M = M_{N_2}$ and the representation at $z = z_1$ is

$$n(z_1, M_{N_2})^{A_{N_2}} = n_d(z_1, M_{N_2})^{A_{N_2}} + n_m(z_1, M_{N_2})^{A_{N_2}}$$
 (A4.1)

where

$$A_{N_2} = M_h / (\overline{M}_0 - M_{N_2}) \tag{A4.2}$$

and $M_h = M_{N_2} = 28$, $M_o = 28.95 \text{ kg(kmol)}^{-1}$. On multiplying Eq. (A4.1) by $[(R_1/A_{N1})T(z_1)]A_{N_2}$, where R_1 and A_{N1} are defined by Eqs. (A.6) and (A73), we obtain a relation in terms of the N_2 pressure $p(z_1, M_{N_2})$ at height z_1

$$p(z_1, M_{N_2})^{A_{N_2}} = p_d(z_1, M_{N_2})^{A_{N_2}} \div p_m(z_1, M_{N_2})^{A_{N_2}}$$
 (A4.3)

where

$$p_d(z, M_i) = (R/A_N) n_d(z, M_i) T(z)$$
 (A4.4)

$$p_{m}(z, M_{i}) = (R/A_{N}) n_{m}(z, M_{i}) T(z)$$
 (A4.5)

with $M_i = M_{N_2}$.

By integration of the hydrostatic equation, the diffusive and mixing pressures are

$$K p_d(z, M_i) = p_d(z_1, M_i) e^{-M_i I(z_1, z)} [T(z_1)/T(z)]^{\alpha} i$$
 (A4.6)

$$K p_{m}(z, M_{i}) = p_{m}(z_{1}, M_{i}) e^{-\overline{M}_{0} I(z_{1}, z)}$$
(A4.7)

where

$$I(z_1, z) = \int_{z_1}^{z} (g/RT) dz$$
(A4.8)

 $K = k(z) = (R_1/A_{N1})/[R(z)/A_N(z)]$.

On dividing Eq. (A4.3) by $\left[p_{m}(z_{2}, M_{N_{2}})\right]^{A_{N_{2}}}$ we obtain, using Eq. (A4.4) to Eq. (A4.7) with $M_{i} = M_{N_{2}}$, $\alpha_{i} = 0$

$$\mu_{N_2} = e^{(\overline{M}_0/M_h)X} (\nu_{N_2}^{A_{N_2}} - e^{-X} + 1)^{A_{N_2}^{-1}}$$
(A4.9)

where

$$X = M_h I(z_1, z_2)$$
 (A4.10)

$$\mu_{N_2} = f_{N_2} p(z_1)/K(z_2) p_m(z_2, M_{N_2})$$
 (A4.11)

$$\nu_{N_2} = n_d(z_2, M_{N_2})/n_m(z_2, M_{N_2})$$
 (A4.12)

 f_{N_2} is the N_2 volume fraction of air at height z_1 and $p(z_1)$ is air pressure at height z_1 , which is evaluated as

$$p(z_1) = p_{ref}(z_1) (1 + D_0) (1 + \Delta_0)$$
 (A4.13)

where p_{ref} is the zonal mean value for a reference value of solar activity (sunspot number) and D_0 , Δ_0 are relative incremental increases associated with longitude and sunspot number respectively. We have

$$D_o = [p(\lambda) - p_{ref}]/p_{ref} = p_1 \cos(\lambda - \lambda_{p1}) + p_2 \cos(2\lambda - \lambda_{p2})$$
 (A4.14)

where p_1 , p_2 , λ_{p1} , λ_{p2} are tabulated in Reference A1. Corresponding to Eq. (A2.22), we have from Reference A3

$$\Delta_{\mathbf{Q}} = \mathbf{K}(\phi) \ell \Phi_{\mathbf{Q}} \Delta \mathbf{R}_{\mathbf{p}} \tag{A4.15}$$

where

$$\Phi_{0} = [1 + \tanh Z_{S} + \frac{1}{4}\alpha (\tanh^{4} Z_{S} - 1)]$$

$$z = z_{1}$$
(A4.16)

The required pressure ratio condition, which is expressed by Eq. (A1.10), is

$$\ell_1 = (M_{z_1}/M_h) X/z_d$$
 (A4.17)

by Eqs. (A1), (A5), (A4.8) and (A4.10)

Appendix A 5 Number Densities

For a gas constituent of molecular weight M_i , the MSIS-86 representation A2 of number density $n(z, M_i)$ introduces factors $C_{1i}(z)$, $C_{2i}(z)$ to account for effects of chemistry and flow:

$$C_{ji}(z) = \exp \left\{ r_{ji} / [1 + \exp(z - z_{ji}) / H_{ji}] \right\}$$
 (j = 1, 2) (A5.1)

 C_{1i} = 1 for N_2 only (i = 3) and C_{2i} = 1 for all gases except O, H and N (i = 2, 7 and 8). The values adopted here for z_{ji} , H_{ji} (j = 1, 2) and r_{2i} are those presented in MSIS-86 for all gases. So also are the values adopted here for r_{1i} (i = 2, 7 and 8) as these gases (O, H and N) are not required to match significant volume fractions of air at height z_1 (= 70 km). The other gases (He, O_2 and Ar) are however required to match significant volume fractions f_i (i = 1, 4 and 5) at height z_1 and this is achieved by means of the relation

$$r_{1i} = \frac{1}{2} \left[(R'_{1i} + R_{1i}) + (R_{1i} - R'_{1i}) \tanh c\zeta \right]$$
 (A5.2)

valid for $z_1 \le z \le z_2$, where c and ζ are defined by Eq. (A4) and Eq. (A5). $R_{1i}^{'}$ depends on f_i and is derived as shown below. Values of $R_{1i}^{'}$ computed in this way are found to lie within about one percent of $R_{1i}^{'}$. The adjustment using Eq. (A5.2) to provide a smooth transition between lower and upper models (for He, O_2 and Ar concentrations) is therefore of negligible physical consequence.

For the calculation of $n(z, M_i)$ we have $^{\mbox{A2}}$

$$n(z, M_i) = \left[n_d(z, M_i)^A_i + n_m(z, M_i)^A_i \right]^{A_i^{-1}} C_{1i}(z) C_{2i}(z) \quad (cm^{-3})$$
 (A5.3)

$$A_{i} = M_{h}/(\overline{M}_{O} - M_{i}) \tag{A5.4}$$

$$n_{d}(z, M_{i}) = n_{d}(z_{2}, M_{i}) e^{M_{i}J(\zeta)} \left[T(z_{2})/T(z)\right]^{1+\alpha_{i}}$$
 (cm⁻³) (A5.5)

$$n_{m}(z, M_{i}) = n_{m}(z_{2}, M_{i}) \overline{M}_{0}^{J(\zeta)} T(z_{2})/T(z)$$
 (em⁻³) (A5.6)

where α_i is the thermal diffusion coefficient and by Eqs. (A4.8), (A1) and (A5)

$$J(\zeta) = I(z, z_2) = [U(1) - U(\zeta)]z_d/M_{z_1}$$
 (A5.7)

where

$$U(\zeta) = \int_{-\infty}^{\zeta} W^{-1} d\zeta . \qquad (A5.8)$$

By Eq. (A1.1) to Eq. (A1.3)

$$U(\zeta) = \sum_{s=1}^{S} \sum_{n=1}^{N} a_{sn} \, \phi_n(\zeta) \, \xi^{s-1} + b_1 \zeta + \frac{1}{2} b_2 \zeta^2 + \dots + \frac{1}{7} b_7 \zeta^7$$
 (A5.9)

$$\phi_{n}(\zeta) = \int_{0}^{\zeta} (1 - \zeta^{2})^{3} (\zeta^{n} - \gamma_{n}) d\zeta = \zeta \left[\psi_{1}(J, n) - \gamma_{n} \psi_{2}(\zeta) \right]$$
(A5.10)

$$\psi_1(\zeta, n) = \zeta^n \left[\frac{1}{n+1} - \frac{3\zeta^2}{n+3} + \frac{3\zeta^4}{n+4} - \frac{\zeta^6}{n+7} \right]$$
 (A5.11)

$$\psi_2(\zeta) = 1 - \zeta^2 + \frac{3}{5}\zeta^4 - \frac{1}{7}\zeta^6 . \tag{A5.12}$$

To evaluate $R_{1i}^{'}$ (i = 1, 4, or 5) we note that C_{2i} = 1 (i = 1, 4 or 5) and put $z = z_1$ in Eq. (A5. 3) to obtain, on multiplying through by $(R_1/A_{N1})T(z_1)/p_m(z_2, M_i)$ and using Eq. (A4. 4) to Eq. (A4. 7) and Eq. (A5. 7),

$$\mu_{i} = \left\{ \begin{bmatrix} \nu_{i} & e^{M_{i} J(-1)} & \left(\frac{T(z_{2})}{T(z_{1})}\right)^{\alpha} \end{bmatrix}^{A_{i}} + \begin{bmatrix} \overline{M}_{0} J(-1) \end{bmatrix}^{A_{i}} \right\}^{A_{i}^{-1}} C_{1i}(z_{1}) \quad (A5.13)$$

where

$$\mu_{i} = f_{i} p(z_{1})/\kappa(z_{2}) p_{m}(z_{2}, M_{i})$$
 (A5.14)

$$\nu_i = n_d(z_2, M_i)/n_m(z_2, M_i)$$
 (A5.15)

By Eqs. (A5.1) and (A5.13), we may solve for r_{1i} (= $R_{1i}^{'}$), where

$$R'_{1i} = \left\{ 1 + \exp\left[(z_1 - z_{1i}) / H_{1i} \right] \right\} \ln C_{1i}(z_1)$$
 (A5.16)

and

$$\ln C_{1i}(z_1) = \ln \mu_i - A_i^{-1} \ln \left\{ \left[\nu_i e^{M_i J(-1)} \left(\frac{T(z_2)}{T(z_1)} \right)^{\alpha_i} \right]^{A_i} + \left[e^{\overline{M}_0 J(-1)} \right]^{A_i} \right\}$$
(A5. 17)

Appendix B

Method of Determination of a sn

We write Eq. (A1.2) as

$$\sum_{s=1}^{S} \sum_{n=1}^{N} C(s, n) a_{sn} = A$$
(B1)

where

$$C(s, n) = (1 - \zeta^2)^3 (\zeta^n - \gamma_n) \xi^{s-1}$$
 (B2)

and determine a_{sn} by the method of weighted least-squares from sets of values $C_k(s,n)$ ($s=1,\ldots,S$; $n=1,\ldots,N$), A_k for $k=1,\ldots,K$. The weighting is based on estimated standard deviations of A_k , σA_k , which are obtained as described below.

The K sets of values are taken at grid points located on a meridional cross-section at each 5 km height interval from z_1 to z_2 (excluding the end points where $\zeta=\pm$ 1 and no contribution is made to the determination of a_{sn}) and at each 10° latitude from pole to pole. Then

$$K = 19[(z_2 - z_1)/5 - 1] . (B3)$$

Likewise $\overline{F}_{10.7}$ is obtained from the three-monthly sunspot number. A_p may be specified for any launch or otherwise set equal to 10. Few profiles extend above 100 km at high latitudes (see Table 3) where dependence on A_p may be important. After examining A_p values for a sample of such launches, the approximation $A_p = 10$ for all launches appeared to be justified.

Dependence on local solar time in the upper model is included, when known, as in the case of single temperature profiles. For monthly mean temperature data, B is calculated from Eq. (A14), that is, Eq. (A1.3), for the diurnal mean upper model with the exception of incoherent scatter data which are obtained during day-time hours only. For I.S. data, B is evaluated for each hour between 0800 and 1600 hours and the average is taken of these 11 values.

The weighted average of the values of A calculated from Eq. (B4) for the k th grid point is based on those temperatures at the height of the k th grid point that lie within 15° latitude of the k th grid point and whose date of observation lies within 1.5 months of the middle of the given month. The weighting factor used in the averaging process is

$$\exp\left\{-\frac{1}{2}\left[\left(\Delta\phi/10\right)^2 + \Delta^2 \text{ m}\right]\right\} \tag{B7}$$

where $\Delta \phi$ is the latitude displacement from the grid point (in degrees) and Δm is the data displacement (in months), $\Delta \phi/10$, Δm being less than 1.5. The most extreme data point to be included in the average is therefore at $\Delta \phi = 15^{\circ}$, $\Delta m = 1.5$ and has a weight of $e^{-2.25}$ (= 0.11) compared with a data point at the grid point ($\Delta \phi = m = 0$) for which the weight is unity.

Temperature data are available either as single observations or as monthly mean values, but the relative accuracies of the various types of data are invariably unknown. The final determination of a sn are not however, sensitive to these relative accuracies as data for single profiles and monthly means are from different locations and tend not to combine.

The introduction of the weighting factor Eq. (B7) results in fractional values for the weighted number of observations, $N_{\rm w}$, and the minimum value for $N_{\rm w}$ below which the available temperature data are considered insufficient for the determination of $A_{\rm k}$ is taken to be 1.1 In that case $A_{\rm k}$ is taken to be the value of $A_{\rm k}$ for the same latitude in the opposite hemisphere with the month shifted by six months, if such a determination is possible with the data available. By this procedure, it was found that $A_{\rm k}$ could be determined at all grid points except for a number at high latitude (60° or more) and above 100 km. For these grid points $A_{\rm k}$ is determined from Eq. (B4) with T equal to its MSIS-86 value.

Likewise $\overline{F}_{10.7}$ is obtained from the three-monthly sunspot number. A_p may be specified for any launch or otherwise set equal to 10. Few profiles extend above 100 km at high latitudes (see Table 3) where dependence on A_p may be important. After examining A_p values for a sample of such launches, the approximation $A_p = 10$ for all launches appeared to be justified.

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$$\exp\left\{-\frac{1}{2}\left[\left(\Delta\phi/10\right)^2 + \Delta^2 \text{ m}\right]\right\} \tag{B7}$$

where $\Delta \phi$ is the latitude displacement from the grid point (in degrees) and Δm is the data displacement (in months), $\Delta \phi/10$, Δm being less than 1.5. The most extreme data point to be included in the average is therefore at $\Delta \phi \approx 15^{\circ}$, $\Delta m = 1.5$ and has a weight of $e^{-2.25}$ (= 0.11) compared with a data point at the grid point ($\Delta \phi = m = 0$) for which the weight is unity.

Temperature data are available either as single observations or as monthly mean values, but the relative accuracies of the various types of data are invariably unknown. The final determination of a are not however, sensitive to these relative accuracies as data for single profiles and monthly means are from different locations and tend not to combine.

The introduction of the weighting factor Eq. (B7) results in fractional values for the weighted number of observations, $N_{\rm w}$, and the minimum value for $N_{\rm w}$ below which the available temperature data are considered insufficient for the determination of $A_{\rm k}$ is taken to be 1.1 In that case $A_{\rm k}$ is taken to be the value of $A_{\rm k}$ for the same latitude in the opposite hemisphere with the month shifted by six months, if such a determination is possible with the data available. By this procedure, it was found that $A_{\rm k}$ could be determined at all grid points except for a number at high latitude (60° or more) and above 100 km. For these grid points $A_{\rm k}$ is determined from Eq. (B4) with T equal to its MSIS-86 value.

At the k th grid point, σA_k , the standard deviation of the weighted average A_k , is estimated from the standard deviation, σ , of the distribution of values of A in the average. The relation adopted is

$$\sigma A_{k} = \sigma/(N_{w} - 1.1)^{1/2}$$
(B8)

in which σ is also an estimated value. In those cases where A_k is found from MSIS-86 temperatures, σA_k is calculated as

$$\sigma A = \sigma \left(M_{z_1} g / RT \right) = \left(M_{z_1} g / RT_c \right) (\sigma T / T_c)$$
(B9)

where T_c is the temperature value from CIRA 1972, Part 1 and

$$\log_{10} (\sigma T) = a (z - 67.5)/42.5 + b$$
 (B10)

where a = 0.519, b = 0.617, (being values based on CIRA 1972 temperature distributions. B1 Equation (B10) gives $\sigma T = 4.4$, 10.3, 24.0 K at 70, 100, 130 km, respectively.

In those cases where σA_k can be found from Eq. (B7), it was desirable to avoid unreasonably small values of σA_k , which may arise at times with selective distributions of data. Hence σA_k is compared with σA and if found to be less than $\sigma A/3$, we take $\sigma A_k = \sigma A$.

The procedure for determining S and N in Eq. (B1) is to take them as small as possible and consistent with a satisfactory least-squares fit to the values of A_k . The values chosen are S = 7, N = 2.

Sets of values of a determined for each month for Case 1 of Section 4.1 are listed in Table B1 and relate to the middle of the month.

B1. COSPAR Working Group 4, (1972) COSPAR International Reference Atmosphere, CIRA 1972, Akademie Verlag, Berlin.

Table B1. Coefficients an Calculated for Case 1

0.301202E+01 -0.715264E+01
-0.959048E+00 -0.261623E+02 0.108858E+02 -0.467195E+02
0.719224E-01 0.158461E+02
0.666923E+01 0.256542E+02
0.397463E+01 0.267372E+02
0.790258E+00 0.140357E+02
0.332649E+01 -0.103331E+01
0.121652E+01 -0.612739E+01
0.225654E+00 -0.457695E+01
0.423461E+01 -0.678040E+01 -0.157533E+02 0.101777E+01 -0.151417E+02 -0.688751E+02
0.401070E+01 -0.684372E+01 -0.331910E+01 0.827888E+00 -0.122259E+02 -0.762859E+02
0.283359E+01 -0.753740E+00 0.531743E+01 0.127062E+02 -0.520754E+01 -0.751404E+01 -0.23351E+02 -0.544138E+01

Appendix C

Average Temperature Deviations from Model Values and Their Mean with Respect to Different Sizes

Comparisons between observed temperatures, T, and model values $T_{\underline{m}}$ are analyzed in terms of the differences

$$T_{d} = T - T_{m} \tag{C1}$$

at each 5 km height interval. For the i th site we obtain the average deviation at any height as

$$x_i = \sum T_d / N_b$$
 (C2)

where N_b is the number of observations available within a selected group of months. Four such groups are considered, namely (a) the three 'winter' months (DJF in the N hemisphere); (b) the six 'equinox' months (MAMSON); (c) the three 'summer' months (JJA in the N hemisphere); and (d) all months. The standard deviation of x_i was estimated in each case as

$$\sigma_{\mathbf{x_i}} = \left[\sum (T_d - x_i)^2 / N_b (N_b - 1) \right]^{1/2}$$
(C3)

for $N_b \ge 2$.

The mean of \mathbf{x}_i with respect to those sites for which a value of \mathbf{x}_i could be determined was obtained as

$$\overline{x} = \sum_{i} w_{i} x_{i} / N_{s}$$
 (C4)

where

$$w_i = N_s (\sigma_{x_i})^{-2} / \sum_i (\sigma_{x_i})^{-2}$$
 (C5)

and $\boldsymbol{N}_{\boldsymbol{S}}$ is the number of such sites.

The standard deviation of \overline{x} was estimated as

$$\sigma_{\overline{\mathbf{x}}} = \sigma_{\overline{\mathbf{D}}}/(N_{\mathbf{s}} - 2)^{1/2} \tag{C6}$$

where σ_D is the standard deviation of the distribution of the values, x_i , about \overline{x} , being estimated (for $N_s > 2$) from

$$\sigma_{\rm D}^2 = \sum_{\rm i} w_{\rm i} (x_{\rm i} - \overline{x})^2 / (N_{\rm s} - 1)$$
 (C7)

Appendix D Coding of the Model Formulation

In the course of this project Fortran programs have been written and tested for each stage of the model formulation and the work has proceeded to completion with the confidence that the formulation was computationally practical and efficient.

Over 50 programs have been developed, some of which are subprograms to the main calculation while others have served only a transient purpose. The latter category includes test programs (that is, main programs for temporary use in the development and testing of the subprograms) and programs for one-of calculations such as the inversion of matrix Eq. (A1.14) and the calculation of Millstone Hill and St. Santin I.S. temperatures from formulas in the references quoted in this report.

The main programs will be described briefly in this Appendix to indicate some of the chief computational stages of the work. Details of the subprograms and input files that are utilized in conjunction with these main programs are given in Appendix E. The particular way in which the computation breaks down into different subprograms reflects the stage-by-stage development of the method. No retrospective consideration has been given at the present time into restructuring the coding. Computational efficiency nevertheless appears to be quite good. All computations have been carried out interactively on EUCLID (the GEC machine at University College London).

D1. DETERMINATION OF a_{sn} (APPENDIX B)

TDIFAVP4

(Temperature Differences Averaged Program 4) calculates A from Eq. (B4) for each observed temperature at 5-km intervals and takes weighted averages to obtain A_k at each grid point of the height-latitudinal cross-section. σA_k is calculated by Eq. (B8). When temperature data are lacking A_k is calculated from Eq. (B4) for heights above 90 km using the temperature value of MSIS-86 and σA_k is used as a flag and set equal to zero.

(Note: the subprograms listed below need to be added to the main program in all cases. Subprograms that are already contained in a main program are not listed, being usually just short routines.)

Subprograms: POLYZ12S, BCZ12S, BCATZ1S, BCATZ2S,

EARTHS, GTS5S, MSISINS, SCATZ1S.

Input Files: TWP12Q, COEFFDQ, RZDATA, THADATA,

POLYTDD, DUMMYAXD.

Output Files: $TDIF(N_1)$, $TDIF(N_1)Q$, $TD(N_1)SD$.

 (N_1) stands for an integer (the run number and is used to identify the outputs corresponding to different inputs). TDIF (N_1) contains A_k and σA_k ; TDIF $(N_1)Q$ is an inspection file for output from the various stages of the computation; and TD (N_1) SD contains the average temperature deviations from the fitted profile, x_i , and their standard deviations, σx_i , for the i th site at each 5-km height interval, the average being taken four times with respect to data falling in each of the groups of months (i) DJF, (ii) MAMSON, (iii) JJA and (iv) all months.

WADJP6

 $(\underline{W}^{-1} \underline{Adj} \underline{ustment \ Program \ 6})$ obtains $\underline{a_{sn}}$ by weighted least-squares solution of Eq. (B1) for each month of the year. Before doing this, it processes $\underline{A_k}$ and $\underline{\sigma}\underline{A_k}$ for each grid point of the height-latitudinal cross-section as described in Appendix B: (1) if $\underline{A_k}$, $\underline{\sigma}\underline{A_k}$ have been calculated from observed temperature data they remain unchanged; (2) if $\underline{\sigma}\underline{A_k} = 0$ (that is, data are lacking), $\underline{A_k}$, $\underline{\sigma}\underline{A_k}$ are taken as the values for the grid point at the same latitude in the opposite hemisphere with a six-month shift of date, unless this also has $\underline{\sigma}\underline{A_k} = 0$ (that is, data are lacking there as well) when, at the original grid point, $\underline{A_k}$ is left unchanged (having been based on an MSIS-86 temperature) and $\underline{\sigma}\underline{A_k}$ is set equal to $\underline{\sigma}\underline{A}$, the value calculated from Eq. (B9); and (3) any $\underline{\sigma}\underline{A_k}$ which is less than $\underline{\sigma}\underline{A}/3$

is considered improbably small and is replaced by σA .

WADJP6 is for Case 1 of Section 4.1

NAG library routine F04ARF. Subprogram:

TDIF(N₁) Input File:

Output Files: OUTRN(N_1), WADJ(N_1)D.

OUTRN(N₁) is an inspection file and WADJ(N₁)D contains a sn.

WADJP7 is WADJP6 modified to deal with Cases 2 and 3 of Sections 4.2

and 4.3.

D2. COMPARISON OF OBSERVATIONS WITH DERIVED MODEL

TDIFAVP4 re-run. TDIFAVP4 is run as above with the following changes:

as before with DUMMYAXD replaced by Input Files:

 $WADJ(N_1)D$.

TDIF(N2), TDIF(N2)Q, TD(N2)SD. Output Files:

Observed temperature data are now compared with the final model, calculated from Eqs. (A1), (A12) and (A13) using the determinations of a_{sn} which are read in from WADJ(N_1)D, whereas in the first use of TDIFAVP4 above a_{sn} were unknown and were read in as zeroes from DUMMYAXD. The purpose of this computation is to obtain $TD(N_2)SD$ for x_i , σx_i for use with AVTDSDP8.

AVTDSDP8 (Average of mean Temperature Deviations and its Standard Deviation Program 8). For each site, the mean temperature deviations x; from the model with respect to data taken in each of four groups of months, namely DJF, MAMSON, JJA and all year, and estimates of their standard deviations ox, are read from TD(N2)SD. The sites are put into three latitude groups, 0-30, 30-50, and 50-90° N or S and their summer, equinox, winter and all-year mean deviations are averaged and the standard deviations of these averages are obtained from Eqs. (C4) to (C7). Results are shown in Tables 4 through 10 for the all-year mean deviations.

> Subprograms: none

Input File: $TD(N_2)SD.$

Output File: $AV(N_2)Q$.

D3. TABULATION OF ATMOSPHERIC PROPERTIES AT HEIGHTS ABOVE 18 km

III IIDIGIIIO IDOVD 10 IIII		
TVALP	for combinin 18 to 70, 70 geophysical p and so on). only and late pressure. F	e Value Program) was written to develop a method g the outputs of the three models for the regions to 130, and above 130 km for given ranges of parameters (height, date, location, solar activity). The program was first written for temperature rextended to provide composition, density and live subprograms are included in this main program the particular outputs when flagged to do so:
	TPDN	for generating output of the height profiles of temperature and log ₁₀ of pressure, density and total number density at given height increments and their third differences with respect to height (to enable the smoothness of these profiles to be examined at 70 and 130 km where continuity in the second height-derivative has been formulated).
	CHPA RM	for generating output of MSIS-86 diffusion and mixing number densities at 130 km, that is, $n_d(z_2, M_i)$ and $n_m(z_2, M_i)$; MSIS-86 density parameter R_{1i} and the amended value R_{1i}' [Eq. (A5.15)]; and turbopause height z_{hi} [Eq. (A96)].
	CHNUMD	for generating output of \log_{10} of the number density of gas constituents, $n(z, M_i)$ at given height increments and their third differences with respect to height (to enable the smoothness of these profiles to be examined at 70 and 130 km where continuity in the second height-derivative has been formulated).
	ATMOS	for generating output of temperature and of \log_{10} of pressure, density, total number density and individual gas number densities.
	ATMOSN	for generating output of temperature, number density, total number, pressure and density.

Subprograms:

TEMPS, POLYZ12S, BCZ12S, BCATZZS,

GTS5S, CF1880S, CFZ1Z2S.

BCATZ2S, SCATZ1S, EARTHS, MSISINS, LINES,

Input Files: TWP12Q, COEFFDQ, WADJ(N₁)D, PARAMD.

Output File: TABLEQ.

TVALIKMP (TVALP amended to give 1 km interval table Program). TPDN and CHNUMD are deleted from TVALP and changes are introduced in conjunction with three subprograms that are added:

TPDTB1 for generating output of separate tables of

temperature, pressure, density and geostiophic W-E wind on each page of output with the same format and an integer height interval (say 1 km) as in the corresponding tables in the report 'A Global Reference Atmosphere From 18 to 80

km'. $^{\Lambda 1}$

POWER for use with TPDTB1 to calculate the power of

10 needed in the pressure and density tables.

GEOW for use with TPDTB1 to calculate geostrophic

W-E wind velocities.

Subprograms and Input Files: as TVAL (but PRM86D1 replaces

PARAMD).

Output File: TAB86Q1.

TAVL5KMP (TAVLP amended to give 5 km interval table Program).

TVALIKMP has subprogram TPDTB1 replaced by TPDTB5 and a few associated changes:

TPDTB5 for generating tables of temperature, pressure,

density and geostrophic wind with an integer height interval (say 5 km) for each of four selected months such that these 12 tables fit into one page of output as shown in Appendix F. Six tables for just two months appear on one page if the number of heights tabulated exceeds

a specified number.

Subprograms and Input Files: as TVAL (but PRM86D5

replaces PARAMD).

Output File: TAB86Q5

Appendix E

Memo List of Subprograms and Input Datafiles

 $COEFFDQ \qquad (\underline{Coefficients\ from\ part\ \underline{D}\ of\ an\ earlier\ \underline{Q}\ (standing\ for\ output))}$

is a datafile containing the coefficients c_{ns} for the lower

model (Reference A1, pp. 108-109).

TPW12Q (Temperature, Pressure Waves $\underline{1}$ and $\underline{2}$ from an earlier \underline{Q}) is a

datafile containing the tables of amplitudes and phases of Reference A1, pp. 110-121, which define the longitudinal dependences of temperature, pressure and density for the lower model. (Only the temperature and pressure dependences

are utilized.)

BCATZ1S (Boundary Conditions at height z_1 Subprogram) is

SUBROUTINE BCATZ1(ZZ, DAY, MN, GLAT, GLONG, G)

which is used with TDIFAVP4 (Appendix D) and interpolates coefficients c_{ns} of the 18 to 80 km region to the given day, DAY, of month, MN, and calculates the required conditions at latitude, GLAT, and height, ZZ (= z_1 = 70 km) transferring these values back to the calling program as G(1), ... $G(4) = p(z_1, M_{N_2})$, 100 W^{-1} and the first two derivatives of 100 W^{-1} with respect to height at height z_1 according to the relations in Appendix A2. Longitude (GLONG) dependence is included unless ILONG1 is set equal to zero in the main program.

BCATZZS (Boundary Conditions at height z₁ or at lower height z Subprogram) is

SUBROUTINE BCATZZ(ZZ, DAY, MN, GLAT, GLONG, G)

which is used with TVALP, TVAL1KMP, TVAL5KMP (Appendix D) and is BCATZ1S(ZZ, DAY, MN, GLAT, GLONG, G) with an additional section of instructions such that when ZZ = 70 (km) it provides the same G(1), ... G(4) as BCATZ1S and for other ZZ (= z) it provides G(1) = p(z, M_{N_2}) and G(2) = 100 W^{-1} at height z (less than 70 km). is the MSIS-86/CIRA 1986 neutral atmosphere model of

GTS5S is the MSIS-86/CIRA 1986 neutral atmosphere model of 15 March 1986 by A.E. Hedin as

SUBROUTINE GTS5(IYD, SEC, ALT, GLAC, GLONG, S1L, F107A, F107, AP, MASS, D, T)

with the additional facility to transfer parameters to other programs through the common block AA defined by

COMMON/AA/ DDF(8), DMX(8), HC04, HC16, BLK1, HC32, HC40, BLK2, HC01, HC14, ZC04, ZC16, BLK3, ZC32, ZC40, BLK4, ZC01, ZC14, BLK5, RC16, BLK6, BLK7, BLK8, BLK9, RC01, RC14, BLK10, HCC16, BLK11, BLK12, BLK13, BLK14, HCC01, HCC14, BLK15, ZCC16, BLK16, BLK17, BLK18, BLK19, ZCC01, ZCC14, RCU(8)

where DDF(I), DMX(I) are respectively the diffusion profile number density $n_{d}(z, M_{I})$ and the mixing profile number density $n_{m}(z, M_{I})$ corresponding to I=1, 2, 3, 4, 5, 7, 8. RCU(I) are likewise the values of R_{1i} (Appendix A5).

BCATZ2S (Boundary Conditions at height z_2 Subprogram) is

70

SUBROUTINE BCATZ2(IYD, SEC, Z2, GLAT, GLONG, STL, F107A, F107, AP, G)

which calls GTS5 to calculate the MSIS-86 conditions required at height $\mathbf{z_2}$ (= 130 km) and transfer them to the calling program as G(1), ... $G(6) = \mathbf{p_m}(\mathbf{z_2}, \mathbf{M_{N_2}})$, $\mathbf{p_d}(\mathbf{z_2}, \mathbf{M_{N_2}})$, $100\,\mathrm{W}^{-1}$ and the first three derivatives of $100\,\mathrm{W}^{-1}$ with respect to height at height $\mathbf{z_2}$ according to the relations in Appendix A3. (G(6) is not utilized.)

BCZ 12S (Boundary Conditions at \underline{z}_1 and \underline{z}_2 Subprogram) is

SUBROUTINE BCZ12(Z1, Z2, IDAY, MN, GLAT, GLONG, SLT, F107A, F107, AP, G1, G2)

and is the calling program of (1) either BCATZ1 (if used with TDIFAVP4) or BCATZZ (if used with TVALP, TVAL1KMP or TVAL5KMP) to obtain G(I) as G1(I) to which the sunspot number dependent changes are applied by calling SCATZ1 and (2) BCATZ2 to obtain G(I) as G2(I).

SCATZ1S (Solar Cycle at height z_1 Subprogram) is

SUBROUTINE SCATZ 1(ZZ, IYD, GLAT, RNDEL, GDEL)

which calculates incremental values GDEL(I) arising from an incremental change of sunspot number RNDEL according to the relations Eqs. (A2. 22) and (A4.15).

RZDATA (Sunspot number, R_z, <u>Datafile</u>) contains monthly mean sunspot numbers from January 1957 to January 1972, the interval of time within which the launch dates of single rocket profiles fail.

POLYZ12S (Polynomial coefficients for height range z_1 to z_2 Subprogram) is

SUBROUTINE POLYZ12(Z1, Z2, IDAY, MN, GLAT, GLONG, SLT, F107A, F107, AP, B)

which calculates the polynomial coefficients \underline{b} from Eq. (A1.15) having first obtained $\underline{\ell}$ which involves the relations of Appendix A4 to obtain ℓ_1 from Eq. (A4.17).

EARTHS (Earth radius and gravity Subprogram) is

SUBROUTINE EARTH(ALAT, GEPHI, RPHI)

which calculates g_{ϕ} and r_{ϕ} from Eqs. (A8) and (A9) at latitude ALAT (= ϕ).

MSISINS (MSIS variations included Subprogram) is

SUBROUTINE MSISIN

which lists those of the 23 variations of MSIS-86 that are being omitted when the Subroutine GTS5 is called.

THADATA (Temperature high altitude Datafile) contains observed temperatures at 5-km height intervals from 70 to 130 km at different sites.

POLYTDD

(Specifies polynomial for obtaining temperature differences Datafile) is a file of input parameters for TDIFAVP4: IH1, IH2 = the range of data in THADATA to be utilized, NOBS = the number of single profiles to which a monthly mean is equated for purposes of weighting (arbitrarily chosen); ISLT = 0 or 1 according to whether local solar time is excluded or included in MSIS-86, ILONG1 = ILONG2 = 0 or 1 according to whether longitude dependences are excluded or included in lower model or upper model (that is, MSIS-86), NCOND = number of coefficients (= 7) of the interpolating polynomial between heights z_1 and z_2 , NF107 (= 120) and NP (= 10) are values of solar activity parameters F_{10.7} and A_p which are adopted when no other values are specified, NRUN = reference number assigned to a particular computer run, IMSIS = 0 for normal use of TDIFAVP4 and = 99 to give tables of T-TMSIS where T is a polynomial model temperature and TMSIS is the corresponding MSIS-86 value, ISAX (= 7) and INAX (= 2) are the values of S and N in Appendix B, NOSCZ1 = 0 for RNDEL to be set zero in BCZ12 so that no sunspot number dependence is introduced into the lower model.

DUMMYAXD

($\underline{\underline{Dummy}}$ set of $\underline{\underline{a}}_{sn}$ coefficients $\underline{\underline{D}}$ atafile) is datafile of $\underline{\underline{a}}_{sn}$ all of which are zero.

TEMPS

(Temperature etc Subprogram) is

SUBROUTINE TEMP(ZHT1, ZHT2, ZHTD, IHEND, Z1, Z2, IDAY, MN, GLAT, GLONG, SLT, F107A, F107, AP, TT, DN, PR, TOTND, RCL, ZZTURB, DDFZ2, DMXZ2)

which calculates at each of IHEND heights from ZHT1 to ZHT2 at a height interval of ZHTD the temperature TT, the number densities of individual gas constituents DN(I), I \neq 6 and the total mass density DN(6), the total number density TOTND, the parameter RCL(I) [= $R_{1I}^{'}$ from Eq. (A5. 16)], the 'turbopause' heights ZZTURB(I) [= z_{hI} of Eq. (A97)], diffusion profile number densities DDFZ2(I) (= $n_d(z_2, M_I)$) and mixing profile number densities DMXZ2(I) (= $n_m(z_2, M_I)$) at height z_2 of the gas constituents.

LINES

(Line Subprogram) is

SUBROUTINE LINE(CHAR)

which reads in and writes out a line of characters.

CF1880S (Coefficients for 18 - 80 km model Subprogram) is
SUBROUTINE CF1880(CC, DELG, WAV)

which reads in TPW12Q and derives DELG and WAV for use in BCATZZS and then reads in COEFFDQ into CC.

CFZ1Z2S (Coefficients a_{sn} for $z_{\underline{1}}$ to $z_{\underline{2}}$ model Subprogram) is

SUBROUTINE CFZ1Z2

PRM86D1

which reads in and writes out a sn.

PARAMD

(Parameter Datafile) runs TVALP with ZHT1, ZHT2, ZHTD;
MN1, MN2, MND; LAT1, LAT2, LATD; LNG1, LNG2, LNGD;
ISLT1, ISLT2, ISLTD as the ranges of height, month,
latitude, longitude and local solar time and their respective
increments at which values are required to be evaluated;
IXLONG = 0 gives zonal mean values; ICHPAR, ITPDN, ICHNUM,
IATMOS, ITMOSN = 0, 1 for output subroutines CLPARM,
TPDN, CHNUM, ATMOS and ATMOSN not to be called or to

called; IXSLT = 0 gives diurnal mean values in MSIS-86;
solar activity parameters for which TVALP is to run are listed
as F107A, F107, AP(1), ... AP(7), SW(9), and NOSCZ1, where
SW(9) controls the use of AP(1) in GTS5 (that is, MSIS-86) and
NOSCZ1 = 0, 1 excludes or includes solar cycle dependence in
the lower model according to SCATZ1.

(Parameters for MSIS-86 Datafile with TVAL1KMP) gives

ZHT1(= 65.), ZHT2 (= 135.), ZHTD (= 1.); LAT1 (= -80);

LAT2 (= 80); LATD (= 10); LNG1, LNG2, LNGD; ISLT1,

ISLT2, ISLTD as the ranges of height, latitude, longitude and local solar time and their respective increments at which values are required to be evaluated; IXLONG = 0 gives zonal mean values; IXSLT = 0 gives diurnal mean values in MSIS-86,

ICHPAR (= 0), IATMOS (= 0); ITMOSN (= 0), ITPDTB (= 1) for output subroutines CHPARM, ATMOS, ATMOSN, TPDTB1 not to be called (= 0) or to be called (= 1); IMSIS = 0 for normal use, otherwise tables of values of MSIS-86 are generated; tables are generated for given dates specified by day of month and month as IDAT(I), MON(I), I = 1, ... NDAYS: IDAT(I) = 0 gives mid-month values; IXMN = 0 gives annual mean, = 1 for given dates as specified above; solar activity parameters for which

TVALIKMP is run and their controlling parameters SW(9) and

NOSCZ1 are supplied as in PARAMD.

PRM86D5

(Parameters for MSIS-86 Datafile with TVAL5KMP) is the same as PRM86D1 with different height and latitude ranges; ZHT1 (= 70.), ZHT2 (= 130.), ZHTD (= 5.); LAT1 (= -80), LAT2 (= 80), LATD (= 20). Geostrophic W-E wind is not tabulated if IGEOTB = 0.

Appendix F

Tabulations of Temperatures, Pressures and Densities, 70-130 km

Diurnal and zonal means of mid-month values of temperature, pressure and density are presented as height-latitude cross-sections for:

- (i) A_p = 4, and 132 and $F_{10.7}$ = 70 and 150 units corresponding to low and medium levels of solar activity. At intermediate levels of solar activity the tables may be interpolated linearly with respect to $F_{10.7}$ and $\log A_p$.
- (ii) A height interval of 5 km from 70 to 130 km. At intermediate heights, pressures and densities may be obtained by interpolation of their logarithms.
- (iii) Case 1 of Section 4.1 using the coefficients in Table B1.

The tables match and are continuous with those of Reference A1 at 70 km and of Reference A2 (MSIS-86) at 130 km.

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	226	218	210	200	190	186	189	203	232	279	343	10	460		3.89	1.83	0.84	3.71	1.58	6.58	2.75	1.21	5.80	3.15	1.93	1.30	9.35		9.00	2.92	1.39	6.47	2.89	1.23	4.99	2.00	8.26	3.65	1.78	0.99	6.21
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<u></u>	227	219	213	203	193	187	191	209	246	302	364	416	424	<u>-</u>	3.02	1.43	99.0	2.95	1.27	5.33	2.24	8	4.96	2.81	1.78	1.22	8.70	<u>.</u>	0.46	2.27	1.08	0.50	2.30	0.99	4.02	1.61	6.67	3.02	1.56	0.92	5.84
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RATURE	215	191	172	162	162	170	189	219	260	31	369	427	481	SURE (N/N SQ)	7.16	3.13	1.24	.52	1.62	5.94	2.37	1.06	5.4	3.09	1.95	1.32	9.45	-	1.16	5.73	2.52	0.98	3.43	1.21	4.30	1.65	9.6	3.25	1.71	9.89	5.19
RATURE	209 215 ;	194 191	184 172	178 162	176 162	179 170	190 189	211 219	244 260	294 311	356 359	419 427	471 481	PRESSURE (N/M SQ)	5.75 7.16	2.50 3.13	1.03 1.24	4.11 4.55	1.61 1.62	6.32 5.94	2.60 2.37	1.15 1.06	5.65 5.41	3.13 3.09	1.93 1.95	1.30 1.32	9.32 9.45	-	0.96 1.16	4.50 5.73	1.95 2.52	0.80 0.98	3.17 3.48	1.22 1.21	4.58 4.30	1.84 1.65	7.68 5.94	3.48 3.25	1.75 1.71	9.84 9.85	6.17 6.19
RATURE	211 209 215	200 194 191	195 184 172	190 178 162	183 176 162	179 179 170	182 190 189	198 211 219	230 244 260	284 294 311	353 326 369	420 419 427	471 471 481	PRESSURE (N/M SQ)	5.31 5.75 7.16	2.35 2.50 3.13	1.01 1.03 1.24	4.28 4.11 4.55	1.76 1.61 1.62	7.07 6.32 5.94	2.85 2.60 2.37	1.21 1.15 1.06	5.68 5.65 5.41	3.06 3.13 3.09	1.88 1.93 1.95	1.27 1.30 1.32	9.11 9.32 9.45	-	0.87 0.96 1.16	4.09 4.50 5.73	1.81 1.95 2.52	0.78 0.80 0.98	3.34 3.17 3.48	1.37 1.22 1.21	5.35 4.58 4.30	2.04 1.84 1.65	8.17 7.68 5.94	3.51 3.48 3.25	1.70 1.75 1.71	9.52 9.84 9.85	5.93 6.17 6.19
RATURE	211 211 209 215 ;	197 200 194 191	191 195 184 172	196 190 178 162	182 183 176 162	180 179 179 170	184 182 190 189	200 198 211 219	230 230 244 260 3	280 284 294 311	347 353 356 369	416 420 419 427	469 471 471 481	PRESSURE (N/H SQ)	5.29 5.31 5.75 7.16	2.34 2.35 2.50 3.13	0.79 1.01 1.03 1.24	4.10 4.28 4.11 4.55	1.55 1.76 1.61 1.62	6.59 7.07 6.32 5.94	2.72 2.85 2.60 2.37	1.17 1.21 1.15 1.06	5.54 5.68 5.65 5.41	2.79 3.06 3.13 3.09	1.83 1.88 1.93 1.95	1.23 1.27 1.30 1.32	8.85 9.11 9.32 9.45	-	0.37 0.87 0.96 1.16	4.12 4.09 4.50 5.73	1.91 1.81 1.95 2.52	0.77 0.78 0.80 0.98	3.18 3.34 3.17 3.48	1.29 1.37 1.22 1.21	5.05 5.35 4.58 4.30	1.97 2.06 1.84 1.65	7.96 8.17 7.68 5.94	3.46 3.51 3.48 3.25	1.67 1.70 1.75 1.71	9.26 9.52 9.84 9.85	5.80 5.93 6.17 6.19
RATURE	211 211 211 209 215	200 197 200 194 191	194 191 195 184 172	187 186 190 178 162	180 182 183 176 162	176 180 179 179 170	180 184 182 190 189	197 200 198 211 219	231 230 230 244 260 3	284 280 284 294 311 ;	350 347 353 356 369	410 416 420 419 427	456 469 471 471 481	PRESSURE (N/M SQ)	5.15 5.29 5.31 5.75 7.16	2.28 2.34 2.35 2.50 3.13	0.98 0.79 1.01 1.03 1.24	4.10 4.10 4.28 4.11 4.55	1.66 :.55 1.76 1.61 1.62	6.60 6.59 7.07 6.32 5.94	2.63 2.72 2.85 2.60 2.37	1.12 1.17 1.21 1.15 1.06	5.28 5.54 5.68 5.65 5.41	2.87 2.79 3.06 3.13 3.09	1.78 1.83 1.88 1.93 1.95	1,20 1,23 1,27 1,30 1,32	8.58 8.85 9.11 9.32 9.45	-	0.37 0.87 0.96 1.16	3.98 4.12 4.09 4.50 5.73	1.76 1.91 1.81 1.95 2.52	0.76 0.77 0.78 0.80 0.98	3.21 3.18 3.34 3.17 3.49	1.29 1.29 1.37 1.22 1.21	5.00 5.05 5.35 4.58 4.30	1.91 1.97 2.06 1.84 1.65	7.55 7.96 8.17 7.68 5.94	3.27 3.46 3.51 3.48 3.25	1.61 1.67 1.70 1.75 1.71	9.26 9.52 9.84 9.85	5.80 5.93 6.17 6.19
TEMPERATURE	220 211 211 211 209 215 ;	210 200 197 200 194 191	205 194 191 195 184 172 1	199 187 186 190 178 162	191 180 182 183 176 162	184 176 180 179 179 170	184 180 184 182 190 189	195 197 200 198 211 219	221 231 230 230 244 260 ;	266 284 280 284 294 311 ;	329 350 347 353 356 369 4	395 410 416 420 419 427	442 456 469 471 471 481	PRESSURE (N/N SQ)	4.30 5.15 5.29 5.31 5.75 7.16	1.98 2.28 2.34 2.35 2.50 3.13	0.88 0.98 0.79 1.01 1.03 1.24	3.87 4.10 4.10 4.28 4.11 4.55	1.55 1.56 1.55 1.76 1.61 1.62	6.87 6.60 6.59 7.07 6.32 5.94	2.83 2.63 2.72 2.85 2.60 2.37	1.21 1.12 1.17 1.21 1.15 1.06	5.61 5.28 5.54 5.68 5.65 5.41	2.95 2.87 2.99 3.06 3.13 3.09	1.77 1.78 1.83 1.88 1.93 1.95	1.18 1.20 1.23 1.27 1.30 1.32	8.33 8.58 8.85 9.11 9.32 9.45	-	0.68 0.95 0.37 0.87 0.96 1.16	3.27 3.98 4.12 4.09 4.50 5.73	1.50 1.76 1.31 1.81 1.95 2.52	0.58 0.76 0.77 0.78 0.80 0.98	3.01 3.21 3.18 3.34 3.17 3.48	1.29 1.29 1.29 1.37 1.22 1.21	5.26 5.90 5.05 5.35 4.58 4.30	2.09 1.91 1.97 2.06 1.84 1.65	8.39 7.55 7.96 8.17 7.68 5.94	3.58 3.27 3.46 3.51 3.48 3.25	1.70 1.61 1.67 1.70 1.75 1.71	9.22 9.10 9.26 9.52 9.84 9.86	5.73 5.75 5.80 5.93 6.17 6.19
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-40 -20 0 20 40 60 80	TEMFERATURE 208 216 220 196 201 197 200 185 184 176 171 189 183 173 167 189 185 179 174 190 192 193 191 200 206 218 224 232 258 276 268 276 317 349 350 558 585	PRESSURE (N/H 5.35 5.44 5.66 5.94 6.38 2.38 2.38 2.54 2.65 2.89 1.03 1.01 1.08 1.11 1.21	4.23 4.46 4.40 4.73 1.77 1.81 1.70 1.77 7.42 7.38 6.64 6.67 3.13 3.10 2.75 2.71 1.37 1.37 1.25 1.23 6.44 6.62 6.37 6.44 3.40 3.57 3.67 3.86 2.05 2.18 2.36 2.58 1.40 1.49 1.66 1.86	DENSITY (KG/M 0-91 0-91 0-94 1-01 4-22 4-41 4-65 4-99 1-83 1-98 2-10 2-30 0-77 0-84 0-87 0-96 3-25 3-44 3-40 3-67 1-36 1-38 1-28 1-33 5-63 5-54 4-89 4-87 2-30 2-25 1-88 9-49 9-43 8-22 7-88 1-91 1-96 1-11 1-17 6-29 6-66 7-24 7-81
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RATURE	218 213 215	205 200 194	1 197 190 177 179	189 182 169	181 178 170	177 179 180	181 190 200	199 214 231	236 256 276	296 321 340	375 403 423	449 478 513	503 534 580 (PRESSURE (N/M S0)	5.46 5.51 6.34 7	2.48 2.46 2.80 3	1.08 1.04 1.14 1	4.58 4.27 4.33 4	1.87 1.70 1.62 1	2 7.45 6.73 6.25 6.29 - 2	2.99 2.77 2.64 2	1.27 1.23 1.24 1	5.08 6.22 6.55 5	3.36 3.58 3.90 3	2.13 2.33 2.57 2	1.49 1.54 1.84 1	1.09 1.22 1.38 1	DENSITY (YG/M CU)	0.97 0.90 1.02 1	4.21 4.28 5.02 5	1 1.91 1.91 2.24 2.56	0.84 0.82 0.89 1	3.59 3.32 3.32 3	1.16 1.30 1.21 1	5.54 5.01 4.54 4	2.15 1.96 1 83 1	8.53 3.10 3.00 \$	3.58 3.57 3.93 3	1.80 1.87 2.01 1	1.03 1.10 1.18 1	6.65 7.23 7.78 8
	219 215 ;	212 204	207 197 191	198 190	188 182	182 178	185 182	202 199	239 233	302 289	185 364	462 439	512 494		5.28	2.38	1.04	4.41	1.81	7.10 7.25 7.32	2.94	1.26	6.01	3.30	2.07	1.43	1.05		0.86	4.05	1.55 1.84 1.91	0.81	3.45	7	5.51	2.13	9.52	3.59	1.89	1.01	6.51
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RATURE	212 215 221 229 3	199 205 213 221	193 199 210 213	190 193 204 207	189 184 195 204	190 179 189 204	196 183 190 208	209 202 205 220	254 240 240 240	278 302 294 295	429 443 459 480	493 493 509 547	PRESSURE (N/M SQ)	4.46 3.68	2.44 2.43 2.07 1.75 1	1.04 1.07 0.94 0.81 (4.35 4.57 4.21 3.66	1.81 1.90 1.84 1.63	7.61 7.68 7.80 7.30	3.27 3.12 3.30 3.32	1.48 1.35 1.46 1.57	7.20 5.58 7.18 7.93 7	3.91 3.71 4.03 4.42	2.40 2.38 2.58 2.78 3	1.63 1.67 1.82 1.95 3	1.23 1.35 1.46 1	DENSITY (KG/M CU)	0.87 0.70 0.56 (4.26 4.14 3.37 2.76 3	1.98 1.96 1.56 1.32 1	0.80 0.82 0.72 0.62 0	3.34 3.58 3.27 2.78 2	1 47 1 58 1 18 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	2.37 2.24 2.40 2.42 1	1.01 0.90 0.99 1.08	4.52 3.94 4.35 4.97 4	2.18 1.93 2.15 2.41	1.18 1.15 1.23 1.30	7.55 7.54 8.15 8.46
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DIVENAL ARP ZORAL MEAN OF MID-NOATH VALUES

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DIUFRAL AND ZOMAL MEAN OF MID-ECATH VALUES

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RATURE	215	161	172	162	162		16	122	266	323	390	462	533	E (N/M SR)			1.24 1.51										1.08 1.09	(KG/M CU)	1.16	5.73		80.0	3.48		4.31	1.65	2.00	3.30	1.73	5.0	6.36
RATURE	215	161	•	162	162		16	122	266	323	390	462	533	SSURE (N/M SO)	7.16	3.13		٠. در	1.62	5.98	2.40	1.09	5.61	73.0	2.10		.08	VSITY (KG/M CU)	1.16	5.73		80.0	3.48		4.31	1.65	2.00	3.30	-	5.0	6.36
RATURE	209 215 3	194 191	172	178 162	177 162	180 171	192 191	214 222	250 266	304 323	375 390	453 462	522 533	PRESSURE (N/M SO)	5.75 7.16	2.50 3.13	1.24	4.11 4.55	1.61 1.62	6.36 5.98	2.63 2.40	1.18 1.09	5.87 5.61	77.0 10.0	2.10 2.10	1.45 1.47	1.07 1.08	DENSITY (KG/H CU)	0.96 1.16	4.50 5.73	1.56 2.52	0.80 0.98	3.17 3.48	1.22 1.21	4.69 4.31	1.85 1.65	7.76 7.00	3.53 3.30	1.73	1.01 1.01	6.36
RATURE	211 209 215	200 194 191	184 172	190 178 162	184 177 162	180 180 171	184 192 191	200 214 222	235 250 266	293 304 323	372 375 390	454 453 462	523 522 533	PRESSURE (N/R SO)	5.75 7.16	2.35 2.50 3.13	1.03 1.24	4.28 4.11 4.55	1.76 1.61 1.62	7.12 6.36 5.98	2.89 2.63 2.40	1.24 1.18 1.09	5.90 5.87 5.61	77.0 10.0 47.0	2.04 2.10 2.10	1.42 1.45 1.47	1.05 1.07 1.08	DENSITY (KG/H CU)	0.87 0.96 1.16	4.09 4.50 5.73	1.81 1.56 2.52	0.78 0.80 0.98	3.33 3.17 3.48	1.37 1.22 1.21	5.37 4.69 4.31	2.07 1.85 1.65	8.27 7.76 7.00	3.57 3.53 3.30	1.78 1.73	0.9/ 1.01 1.01	6.16 6.34 6.36
RATURE	211 211 209 215 ;	198 200 194 191	195 184 172	186 190 178 162	182 184 177 162	181 180 180 171	186 184 192 191	203 200 214 222	235 235 250 266	290 293 304 323 ;	366 372 375 390	450 454 453 462	522 523 522 533	PRESSURE (W/M SQ)	5.75 7.16	2.34 2.35 2.50 3.13	0.99 1.01 1.03 1.24	1.10 4.28 4.11 4.55	1.67 1.76 1.61 1.62	6.73 7.12 6.36 5.98	2.76 2.89 2.63 2.40	1.20 1.24 1.18 1.09	5.76 5.90 5.87 5.61	77.0 10.0 47.0 71.0	1.98 2.04 2.10 2.10	7***	1.02 1.05 1.07 1.08	DENSITY (KG/M CU)	0.87 0.96 1.16	4.12 4.09 4.50 5.73 (1,81 1,81 1,56 2,52	0.73 0.78 0.80 0.98	3.18 3.33 3.17 3.48	1.29 1.37 1.22 1.21	5.07 5.37 4.69 4.31	1.99 2.07 1.85 1.65	8.06 8.27 7.76 7.00	3.52 3.57 3.53 3.30	1.71 1.73 1.78 1.73	0.9/ 1.01 1.01	6.16 6.34 6.36
RATURE	211 211 211 209 215 3	200 198 200 194 191	191 195 184 172	188 186 190 178 162	181 182 184 177 162	171 180 180 171	182 186 184 192 191	260 203 200 214 222	236 235 235 250 266 3	294 290 293 304 323	370 366 372 375 390	444 450 454 453 462	507 522 523 522 533	PRESSURE (N/M SR)	5,15 5,29 5,31 5,75 7,16	2.28 2.34 2.35 2.50 3.13	0.98 0.99 1.01 1.03 1.24	4.10 :.10 4.28 4.11 4.55	1.67 1.67 1.76 1.61 1.62	6.64 6.73 7.12 6.36 5.98	2.67 2.76 2.89 2.63 2.40	1.14 1.20 1.24 1.18 1.09	5.49 5.76 5.90 5.87 5.61	77.5 15.5 47.5 71.5 50.5	1.93 1.98 2.04 2.10 2.10	7*** OF 1 7** 1 20 1 10 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	0.99 1.02 1.05 1.07 1.08	DENSITY (KG/H CU)	0.87 0.96 1.16	3.98 4.12 4.09 4.50 5.73 (1,76 1,81 1,81 1,56 2,52 ;	86.0 08.0 87.0 57.0 87.0	3.20 3.18 3.33 3.17 3.48	1.29 1.29 1.37 1.22 1.21	5.01 5.07 5.37 4.69 4.31	1.92 1.99 2.07 1.85 1.65	7.64 8.06 8.27 7.76 7.00	3.32 3.52 3.57 3.53 3.30	1.64 1.71 1.73 1.78 1.73	0.43 0.45 0.47 1.01 1.01	5.93 5.97 6.16 6.34 6.36
RATURE	220 211 211 211 209 215 ;	210 200 198 200 194 191	194 191 195 184 172	199 188 186 190 178 162	191 181 182 184 177 162	185 177 181 180 180 171	186 182 186 184 192 191	197 260 263 200 214 222	225 236 235 235 250 266	275 294 290 293 304 323 ;	347 370 366 372 375 390	428 444 450 454 453 462	491 507 522 523 522 533	(N/W) BRESSURE (N/W SO)	5,15 5,29 5,31 5,75 7,16	1.98 2.28 2.34 2.35 2.50 3.13	0.88 0.98 0.99 1.01 1.03 1.24	3.87 4.10 1.10 4.28 4.11 4.55	1.66 1.67 1.67 1.76 1.61 1.62	6.91 6.64 6.73 7.12 6.36 5.98	2.87 2.67 2.76 2.89 2.63 2.40	1.24 1.14 1.20 1.24 1.18 1.09	5.82 5.49 5.76 5.90 5.87 5.61	77.0 10.0 7.10 7.10 0.01 0.01 0.01 0.01	1.92 1.93 1.98 2.04 2.10 2.10	7.52 1.54 1.58 1.42 1.45 0.5 0.5 0.5 0.5 0.5 0.5 0.5 0.5 0.5 0.	0.97 0.99 1.02 1.05 1.07 1.08	DENSITY (KG/H CU)	0.87 0.96 1.16	3.27 3.98 4.12 4.09 4.50 5.73 (1.50 1.76 1.81 1.81 1.56 2.52 3	0.68 0.74 0.77 0.78 0.80 0.98	3.01 3.20 3.18 3.33 3.17 3.48	1.29 1.29 1.29 1.37 1.22 1.21	5.27 5.01 5.07 5.37 4.69 4.31	2.10 1.92 1.99 2.07 1.85 1.65	8.48 7.64 8.06 8.27 7.76 7.00	3.64 3.32 3.52 3.57 3.53 3.30	1.73 1.64 1.71 1.73 1.78 1.73	0.94 0.93 0.95 0.97 1.01 1.01	5.92 5.93 5.97 6.16 6.34 6.36
RATURE	227 220 211 211 211 209 215 3	217 210 200 198 260 194 191	205 194 191 195 184 172	267 199 188 186 190 178 162	203 191 181 182 184 177 162	171 081 180 171 181 180 180 171	197 186 182 186 184 192 191	203 197 260 203 200 214 222	222 225 236 235 235 250 266 3	261 275 294 290 293 304 323	330 347 370 366 372 375 390	420 428 444 450 454 453 462	490 491 507 522 523 522 533	PRESSURE (*/# S0)	3.46 4.30 5.15 5.29 5.31 5.75 7.16	1.63 1.98 2.28 2.34 2.35 2.50 3.13	0.98 0.99 1.01 1.03 1.24	3,37 3,87 4,10 1,10 4,28 4,11 4,55	1.50 1.66 1.67 1.67 1.76 1.61 1.62	6.58 6.91 6.64 6.73 7.12 6.36 5.98	2.89 2.87 2.67 2.76 2.89 2.63 2.40	1.29 1.24 1.14 1.20 1.24 1.18 1.09	6.09 5.87 5.49 5.76 5.90 5.87 5.61	77.0 10.0 47.0 71.0 0.00 0.01 0.01	1.91 1.92 1.93 1.98 2.04 2.10 2.10	1.27 1.32 1.34 1.38 1.42 1.40 1.47	0.94 0.97 0.99 1.02 1.05 1.07 1.08	DENSITY (KG/M CU)	0.87 0.96 1.16	3.27 3.98 4.12 4.09 4.50 5.73 (1.23 1.50 1.74 1.81 1.81 1.54 2.52 ;	0.57 0.68 0.74 0.77 0.78 0.80 0.98	2.57 3.01 3.20 3.18 3.33 3.17 3.48	1.15 1.29 1.29 1.29 1.37 1.22 1.21	5.02 5.27 5.01 5.07 5.37 4.69 4.31	2.14 2.10 1.92 1.99 2.07 1.85 1.65	8.48 7.64 8.06 8.27 7.76 7.00	3.93 3.64 3.32 3.52 3.57 3.53 3.30	1.82 1.73 1.64 1.71 1.73 1.78 1.73	0.94 0.93 0.95 0.97 1.01 1.01	5.92 5.93 5.97 6.16 6.34 6.36

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- Ш	221 229	213 221	210 213	205 207	196 204	188 202	139 203	201 211	231 229	284 267	358 334	435 424	197, 496	RESSURE (N/M SQ)	4.46	2.97 1.75	0.74 0.81	4.21 3.66	1.84 1.63	7.79 7.24	3.27 3.25	1.45 1.50	87.7 CB.0	75. 0 25. 0	44 1 50	7.27 7.17		_	9.70 0.29	3.37 2.75	1.55 1.52	0.72 0.61	3.27 2.78	1.43 1.24	5.93 5.45	2.39 2.33	0.97 1.04	4.21 4.56	2.34 2.29	51.1	- 62.5 5.53 5.29 -
- Ш	221 229	213 221	213	205 207	196 204	188 202	139 203	201 211	231 229	284 267	358 334	435 424	197, 496	PRESSURE (N/M SQ)	4.46	2.97 1.75	0.74 0.81	4.21 3.66	1.84 1.63	7.79 7.24	3.27 3.25	1.45 1.50	87.7 CB.0	75. 0 25. 0	44 1 50	· · ·		_	9.70 0.29	3.37 2.75	1.55 1.52	0.72 0.61	3.27 2.78	1.43 1.24	5.93 5.45	2.39 2.33	0.97 1.04	4.21 4.56	3.3	51.1	7.24 7.25 7.33 5.29 -
- Ш	215 221 229 7	205 213 221	210 213	193 205 207	195 196 204	179 188 202	133 139 203	201 201 211	239 231 229	300 284 267	377 358 334	450 435 424	512 497 496	PRESSURE (N/M SQ)	5.39 4.46	2.43 2.97 1.75	1.07 0.74 0.81	4.57 4.21 3.66	1.90 1.84 1.63	7.69 7.79 7.24	3.12 3.27 3.25	1.35 1.43 1.50	2,7 4 0,83 7,78	21 (21 (71 (1 44 1 47 1 50	7.27 7.17		_	0.87 0.70 0.54	3.37 2.75	1.85 1.55 1.52	0.32 0.72 0.81	3.57 3.27 2.78	1.43 1.43 1.24	5.81 5.93 5.45	2.24 2.39 2.33	0.39 0.97 1.01	3.91 4.21 4.56	1.95 2.34 2.29	51.1	024 2.25 7.25 5.29
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SATURE	217 217 212 215 221 229 3	202 203 199 205 213 221 3	191 196 193 199 210 213	184 192 190 193 205 207 3	181 189 189 185 195 204	188 191 179 198 202	193 192 197 133 138 203	213 205 210 201 201 211	247 233 236 239 231 229	300 283 282 300 284 267	373 358 354 377 358 334	454 450 447 450 435 424	524 527 526 512 497, 496	PRESSURE (N/N SD)	6.07 5.75 5.49 5.39 4.46 3.68	2.73 2.50 2.44 2.43 2.97 1.75	1.17 1.13 1.04 1.07 0.74 0.81	4.79 4.78 4.36 4.57 4.21 3.66	1.92 2.00 1.81 1.90 1.84 1.63	7.73 8.35 7.63 7.69 7.79 7.24	3.22 3.53 3.29 3.12 3.27 3.25	00 - 1 - 1 - 1 - 1 - 1 - 1 - 1 - 1 - 1 -	4.03 4.06 3.97 3.68 3.76 3.89	2.54 2.51 2.45 2.36 2.35 2.35	1.77 1.74 1.70 1.66 1.63 1.59	1.31 1.29 1.25 1.24 1.21 1.17	(SITY (0.97 0.93 0.90 0.87 0.70 0.56	4,72 4,45 4,25 4,14 3,37 2,75	2.13 2.00 1.33 1.85 1.55 1.32	0.91 0.34 0.30 0.32 0.72 0.51	3.39 3.47 3.33 3.57 3.27 2.78	1.46 1.53 1.33 1.43 1.43 1.24	5,72 5,28 5,70 5,81 5,93 5,45	2.28 2.55 2.37 2.24 2.38 2.38	3.76 1.35 1.01 0.37 0.97 1.04	3.32 3.62 4.52 3.91 4.21 4.55	Company of the compan	DOLL BUTTON CONTRACTOR STORY
SATURE	221 217 217 212 215 221 229 3	203 202 203 199 205 213 221 3	184 191 196 193 199 210 213 3	172 184 192 190 193 205 207 3	159 181 189 189 185 195 204 3	177 133 188 191 179 198 202	194 193 192 197 193 199 203	222 213 205 210 201 201 211	263 247 233 236 239 231 229	315 300 283 282 300 284 267	393 373 358 354 377 358 334	460 454 450 447 450 435 424	534 524 527 526 512 497, 496	PRESSURE (N/M SQ)	1 6.70 6.07 5.75 5.49 5.39 4.46 3.68	3.04 2.73 2.50 2.44 2.43 2.07 1.75	1.28 1.17 1.13 1.04 1.07 0.74 0.81	5.00 4.79 4.78 4.36 4.57 4.21 3.66	1.87 1.92 2.00 1.81 1.90 1.84 1.63	1 7.16 7.73 8.35 7.63 7.69 7.79 7.24	2.94 3.22 3.53 3.29 3.12 3.27 3.25	00 1 2 3 4 4 4 00 1 2 4 5 1 5 1 5 1 5 1 5 1 5 1 5 1 5 1 5 1 5	2.7 2.9 4.03 4.06 3.97 3.48 3.76 3.89 3.76 3.89 3.89 3.76 3.80 3.76 3.80 3.76 3.80 3.76 3.76 3.76 3.76 3.76 3.76 3.76 3.76	7 2.55 2.54 2.51 2.45 2.36 2.35 2.35	7 1.77 1.77 1.74 1.70 1.56 1.53 1.59	1.31 1.31 1.29 1.25 1.24 1.21 1.17	DENSITY (1.06 0.77 0.93 0.90 0.87 0.70 0.56	5 5.23 4.72 4.45 4.25 4.14 3.37 2.75	5 2.42 2.13 2.00 1.33 1.85 1.55 1.32	1.01 0.71 0.38 0.30 0.42 0.72 0.51	3 3.35 3.39 3.67 3.33 3.57 3.27 2.78	3 1.40 1.16 1.53 1.33 1.43 1.43 1.24	7 5.19 5.72 5.28 5.70 5.81 5.93 5.45	1 2.04 2.28 2.55 2.37 2.24 2.38 2.38	1 3.37 3.76 1.35 1.01 0.37 0.97 1.04	0 1,10 1,32 1,62 1,53 3,91 4,21 4,56 1,0 1,0 1,0 1,0 1,0 1,0 1,0 1,0 1,0 1,0	Control of the contro	COLUMN TO THE CO
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SATURE	225 221 217 217 212 215 221 229 3	204 203 202 203 199 205 213 221 3	184 184 191 196 193 199 210 213 ;	169 172 184 192 190 193 205 207 3	161 169 181 189 189 185 196 204	165 177 133 188 191 179 198 202	182 194 193 192 197 183 188 203	217 222 213 205 210 201 201 211	276 263 247 233 236 239 231 229 3	354 315 300 283 282 300 284 267 3	428 393 373 358 354 377 358 334	183 460 454 450 447 450 435 424	548 534 524 527 526 512 497, 496	EMBER PRESSURE (N/M SQ)	7.24 6.70 6.07 5.75 5.49 5.39 4.46 3.68	3.32 3.04 2.73 2.50 2.44 2.43 2.07 1.75	1.40 1.28 1.17 1.13 1.04 1.07 0.74 0.81	5.41 5.00 4.79 4.78 4.36 4.57 4.21 3.66	1.96 1.87 1.92 2.00 1.81 1.90 1.84 1.63	7.05 7.16 7.73 8.35 7.63 7.69 7.79 7.24	2.70 2.94 3.22 3.53 3.29 3.12 3.27 3.25	00.1 04.1 00.1 74.1 00.1 14.1 40.1 71.1 7	3.59 3.98 4.03 4.06 3.97 3.88 3.78 3.89	2.47 2.55 2.54 2.51 2.45 2.36 2.35 2.35	1.77 1.77 1.77 1.74 1.70 1.56 1.53 1.59	1.31 1.31 1.31 1.29 1.25 1.24 1.21 1.17	DENSITY (1.06 0.77 0.93 0.90 0.87 0.70 0.56	5 5.23 4.72 4.45 4.25 4.14 3.37 2.75	5 2.42 2.13 2.00 1.33 1.85 1.55 1.32	1.01 0.71 0.38 0.30 0.42 0.72 0.51	3 3.35 3.39 3.67 3.33 3.57 3.27 2.78	3 1.40 1.16 1.53 1.33 1.43 1.43 1.24	7 5.19 5.72 5.28 5.70 5.81 5.93 5.45	1 2.04 2.28 2.55 2.37 2.24 2.38 2.38	1 3.37 3.76 1.35 1.01 0.37 0.97 1.04	0 1,10 1,32 1,62 1,53 3,91 4,21 4,56 1,0 1,0 1,0 1,0 1,0 1,0 1,0 1,0 1,0 1,0	Control of the contro	
SATURE	225 221 217 217 212 215 221 229 3	204 203 202 203 199 205 213 221 3	184 184 191 196 193 199 210 213 ;	169 172 184 192 190 193 205 207 3	161 169 181 189 189 185 196 204	177 133 188 191 179 198 202	182 194 193 192 197 183 188 203	217 222 213 205 210 201 201 211	276 263 247 233 236 239 231 229 3	354 315 300 283 282 300 284 267 3	428 393 373 358 354 377 358 334	183 460 454 450 447 450 435 424	548 534 524 527 526 512 497, 496	NOVERBER PRESSURE (N/M SQ)	7.24 6.70 6.07 5.75 5.49 5.39 4.46 3.68	3.32 3.04 2.73 2.50 2.44 2.43 2.07 1.75	1.40 1.28 1.17 1.13 1.04 1.07 0.74 0.81	5.41 5.00 4.79 4.78 4.36 4.57 4.21 3.66	1.96 1.87 1.92 2.00 1.81 1.90 1.84 1.63	7.05 7.16 7.73 8.35 7.63 7.69 7.79 7.24	2.94 3.22 3.53 3.29 3.12 3.27 3.25	00.1 04.1 00.1 74.1 00.1 14.1 40.1 71.1 7	3.59 3.98 4.03 4.06 3.97 3.88 3.78 3.89	2.47 2.55 2.54 2.51 2.45 2.36 2.35 2.35	1.77 1.77 1.77 1.74 1.70 1.56 1.53 1.59	1.31 1.31 1.31 1.29 1.25 1.24 1.21 1.17	DENSITY (1.06 0.77 0.93 0.90 0.87 0.70 0.56	5 5.23 4.72 4.45 4.25 4.14 3.37 2.75	5 2.42 2.13 2.00 1.33 1.85 1.55 1.32	1.01 0.71 0.38 0.30 0.42 0.72 0.51	3 3.35 3.39 3.67 3.33 3.57 3.27 2.78	3 1.40 1.16 1.53 1.33 1.43 1.43 1.24	7 5.19 5.72 5.28 5.70 5.81 5.93 5.45	1 2.04 2.28 2.55 2.37 2.24 2.38 2.38	1 3.37 3.76 1.35 1.01 0.37 0.97 1.04	0 1,10 1,32 1,62 1,53 3,91 4,21 4,56 1,0 1,0 1,0 1,0 1,0 1,0 1,0 1,0 1,0 1,0	Control of the contro	COLUMN TO THE CO

AP =132.0 F107 =150.0

DIURNAL AND ZONAL NEAN OF MID-MONTH VALUES

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80 DEG	(4)	200	221	214	205	197	194	202	224	267	340	442	549		\$0)	3.47 + 0	1.65	0.76		6.48 - 2	2.84	1.33	6.95 - 3	4.13	2.78	2.04	1.57	(n)	0.53 - 4	2.60 - 5	1.24	0.59	2.66 - 6	1.16	4.85 - 7	2.02	0.88	4.05 - 8	2.08	8.20 - 9	
	ATILES (V)		• • •		••	_	_	•	•••	•	.,	•	506 549		(BS H/H)	3.88 3.47 + 0	1.83 1.65	<u> </u>	1,78 1,50	~	•	_	~	•	•••	• •	_	(KG/M CU)	0.60 0.53 - 4	2.89 2.60 - 5	_	_	• •	-	-	•••	_		2.37 2.08		
	(V) Palled (V)	340144	221	21,7	212	207	202	202	211	237	290	383	• •		SSURE (N/M SQ)	80	83		, .	7.97	3.56	1.64	8.07	4.45	2.81	2.00	1.53	SITY (KG/M CU)	0.72 0.60 0.53 - 4	2.89	1.37	0.64	3.00	1.37	90.9	2.62	-13	2.05		8.17	
	TOKERBATHER (V)	TENTERNIONE .	215 221	211 217	206 212 3	198 207	191 202 1	191 202 ;	203 211	233 237 ;	290 290	378 383	908		PRESSURE (N/M SQ)	80	2.11 1.83	0.96 0.85	1 01 1 78	8.23 7.97	3.52 3.56	1.56 1.64	7.57 8.07	4.18 4.45	2.66 2.81	1.89 2.00	1.43 1.53	DENSITY (KG/M CU)	0.72 0.60	3.41 2.89	1.59 1.37	0.73 0.64 (3.36 3.00	1.49 1.37 1	6.28 6.06	2.58 2.62	1.07 1.13 (4.64 5.02	2.21 2.37	7.90 8.17	
	(V) Palleganot	יבחר האוטאני	207 215 221	199 211 217	192 206 212 3	183 198 207	178 191 202 1	183 191 202 ;	202 203 211	241 233 237 ;	305 290 290	390 378 383	480 506		PRESSURE (N/M SQ)	5.43 4.54 3.88	2.48 2.11 1.83	1.09 0.96 0.85	1 97 1 91 1 78	7.76 8.23 7.97	3.15 3.52 3.56	1.36 1.56 1.64	6.64 7.57 8.07	3.76 4.18 4.45	2.44 2.66 2.81	1.74 1.89 2.00	1.31 1.43 1.53	DENSITY (KG/M CU)	0.87 0.72 0.60 (4.17 3.41 2.89	1.90 1.59 1.37	0.85 (.73 0.64 (3.65 3.36 3.00	1.50 1.49 1.37 1	5.87 4.28 6.06	2.26 2.58 2.62	0.90 1.07 1.13 (3.94 4.64 5.02	1.96 2.21 2.37	7.35 7.90 8.17	
	(V) Ballbaganat	CHICKNIONS	203 207 215 221	195 199 211 217	189 192 206 212 3	182 183 198 207 1	179 178 191 202 1	184 183 191 202 3	203 202 203 211 3	241 241 233 237 3	303 305 290 290	386 390 378 383	471 480 506			5.67 5.43 4.54 3.88	2.56 2.48 2.11 1.83	1.11 1.09 0.96 0.85	1.97 1 97 1 91 1.78	7.70 7.76 8.23 7.97	3.13 3.15 3.52 3.56	1.36 1.36 1.56 1.64	6.63 6.64 7.57 8.07 (3.74 3.76 4.18 4.45	2.41 2.44 2.66 2.81	1.71 1.74 1.89 2.00	1.28 1.31 1.43 1.53	DENSITY (KG/M CU)	0.90 0.87 0.72 0.60 (4.39 4.17 3.41 2.89	1.98 1.90 1.59 1.37	0.86 0.85 0.73 0.64 (3.65 3.65 3.36 3.00	1.48 1.50 1.49 1.37 1	5.80 5.87 4.28 4.06	2.24 2.26 2.58 2.62	0.90 0.90 1.07 1.13 (3.95 3.94 4.64 5.02	1.96 1.96 2.21 2.37	7.35 7.90 8.17	
	(A) Belled Bangt	THE VEC ALC BIC ALC	204 203 207 215 221	198 195 199 211 217	193 189 192 206 212 3	186 182 183 198 207 1	182 179 178 191 202 1	186 184 183 191 202	205 203 202 203 211 ;	243 241 241 233 237 3	307 303 305 290 290	394 386 390 378 383	519 577 550 506			5.67 5.43 4.54 3.88	2.55 2.56 2.48 2.11 1.83	1.11 1.11 1.09 0.96 0.85 C	1.08 1.90 1.07 1.01 1.78 1	8.08 7.70 7.76 8.23 7.97	3.31 3.13 3.15 3.52 3.56	1.44 1.36 1.36 1.56 1.64	7.05 6.63 6.64 7.57 8.07 (3.97 3.74 3.76 4.18 4.45	2.56 2.41 2.44 2.66 2.81	1.82 1.71 1.74 1.89 2.00	1.37 1.28 1.31 1.43 1.53	DENSITY (KG/M CU)	0.91 0.90 0.87 0.72 0.60	4.36 4.39 4.17 3.41 2.89	1.96 1.98 1.90 1.59 1.37	0.86 0.86 0.85 0.73 0.64 (3.71 3.65 3.65 3.36 3.00	1.54 1.48 1.50 1.49 1.37 1	6.08 5.80 5.87 6.28 6.06	2.37 2.24 2.26 2.58 2.62	0.96 0.90 0.90 1.07 1.13 (4.17 3.95 3.94 4.64 5.02	2.06 1.96 1.96 2.21 2.37	7.61 7.19 7.35 7.90 8.17	
	(V) Ballbagawat	THE VEG OIC TIC FIC	198 204 203 207 215 221 3	187 198 195 199 211 217	181 193 189 192 206 212 3	179 186 182 183 198 207 1	184 182 179 178 191 202 1	196 186 184 183 191 202 3	221 205 203 202 203 211 3	262 243 241 241 233 237 3	326 307 303 305 290 290	414 394 386 390 378 383	482 469 471 480 506 551 519 519 550 5			6.24 5.66 5.67 5.43 4.54 3.88	2.76 2.55 2.56 2.48 2.11 1.83	1.16 1.11 1.11 1.09 0.96 0.85	1.07 1.08 1.00 1.07 1.04 5.75 1.05 1.05 1.05 1.05 1.05 1.05 1.05 1.0	7,49 8.08 7.70 7.76 8.23 7.97	3.16 3.31 3.13 3.15 3.52 3.56	1.45 1.44 1.36 1.36 1.56 1.64 1	7.42 7.05 6.63 6.64 7.57 8.07 (4.31 3.97 3.74 3.76 4.18 4.45	2.82 2.56 2.41 2.44 2.66 2.81	2.02 1.82 1.71 1.74 1.89 2.00	1.53 1.37 1.28 1.31 1.43 1.53	DENSITY (KG/M CU)	1.02 0.91 0.90 0.87 0.72 0.60 (4.86 4.36 4.39 4.17 3.41 2.89	2.16 1.96 1.98 1.90 1.59 1.37	0.90 0.86 0.86 0.85 0.73 0.64	3.61 3.71 3.65 3.65 3.36 3.00	1.41 1.54 1.48 1.50 1.49 1.37	5.51 6.08 5.80 5.87 6.28 6.06	2.22 2.37 2.24 2.26 2.58 2.62	0.94 0.96 0.90 0.90 1.07 1.13 (4.33 4.17 3.95 3.94 4.64 5.02	2.20 2.06 1.96 1.96 2.21 2.37	8.24 7.61 7.19 7.35 7.90 8.17	
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INPERATURE (K)	215	161	172	163	163	174) i	• • •	370	462	554	632	SSURE (N/M SO)	7.16	3.14	1.2	1.64	6.10	2.51		1.92	2.57	1.96	1.50		1.15	5.73	2.52	0.98	3.47	1.21	38	2,70	7.32		2 10	7.73
TEMPERATURE (K)	209 215	194 191	184 172	178 163	177 163	182 174	196 197	293	331 370	418 462	510 554	585 632	PRESSURE (N/M SQ)	5.75 7.16	2.50 3.14	1.03 1.24	1.62 1.64	6.44 6.10	2.70 2.51	1.24 1.18	3.77 3.92	2.45 2.57	1.75 1.96	1.33 1.50		0.96 1.15	4.50 5.73	1.94 2.52	0.80 0.78	3.17 3.49	1.22 1.21	4.73 4.38	1.39 1.70	3 40 7 52	1 20 - 00 -	1.10 1.15	7.18 7.73
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TEMPERATURE (K)	211 211 209 215	198 200 194 191	191 195 184 172	186 190 178 163	183 184 177 163	181 181 182 174	761 661 781 781	239 241 265 293	296 305 331 370	377 392 418 462	465 481 510 554	539 553 585 632	PRESSURE (N/M SD)	5.29 5.31 5.75 7.16	2.34 2.35 2.50 3.14	0.99 1.01 1.03 1.24	1.67 1.77 1.62 1.64	6.76 7.16 6.44 6.10	2.79 2.93 2.70 2.51	1.22 1.27 1.24 1.18	3.30 3.47 3.72 3.92	2.10 2.23 2.45 2.57	1.48 1.58 1.75 1.94	1.11 1.19 1.33 1.50		0.87 0.87 0.94 1.15	4.12 4.09 4.50 5.73	1.81 1.91 1.94 2.52	0.77 0.78 0.80 0.78	3.18 3.33 3.17 3.49	1.29 1.37 1.22 1.21	5.08 5.40 4.73 4.38	2.00 2.10 1.39 1.70	3.16 3.44 3.91 /.32 6	20 1 20 1 20 1 20 1 20 1 20 1 20 1 20 1	0.93 1.03 1.10 1.15	5.24 5.62 7.18 7.73
TEMPERATURE (K)	211 211 211 209 215	200 198 200 194 191	194 191 195 184 172	183 186 190 178 163	181 183 184 177 163	178 181 181 182 174	183 187 183 196 197	263 204 204 222 230 3	306 296 305 331 370	390 377 392 418 462	471 465 481 510 554	537 539 553 585 632	PRESSURE (N/M SQ)	5.15 5.29 5.31 5.75 7.16	2.28 2.34 2.35 2.50 3.14	0.98 0.99 1.01 1.03 1.24	1.67 1.67 1.77 1.62 1.64	6.67 6.76 7.16 6.44 6.10	2.70 2.79 2.93 2.70 2.51	1.17 1.22 1.27 1.24 1.18	3.26 3.30 3.47 3.72 3.92	2.11 2.10 2.23 2.45 2.57	1.50 1.48 1.58 1.75 1.96	1.13 1.11 1.19 1.33 1.50		0.85 0.87 0.87 0.94 1.15	3.98 4.12 4.09 4.50 5.73	1.76 1.81 1.81 1.94 2.52	0.75 0.77 0.78 0.80 0.98	3.20 3.18 3.33 3.17 3.49	1.30 1.29 1.37 1.22 1.21	5.03 5.08 5.40 4.73 4.38	1.94 2.00 2.10 1.39 1.70	7 4 4 4 4 60 4 60 4 60 4 60 4 60 4 60 4		0.98 0.98 1.03 1.04 1.15	6.37 5.24 5.62 7.18 7.73
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TEMPERATURE (K)	227 220 211 211 211 209 215	217 210 200 198 200 194 191	211 204 194 191 195 184 172	207 199 188 186 190 178 163	204 192 181 183 184 177 163	201 187 178 181 181 182 174	203 189 183 187 185 196 197	238 242 239 241 265 293	294 297 306 296 305 331 370	386 386 390 377 392 418 462	506 484 471 465 481 510 554	589 552 537 539 553 585 632	PRESSURE (N/M SD)	3.46 4.30 5.15 5.29 5.31 5.75 7.16	1.63 1.98 2.28 2.34 2.35 2.50 3.14	0.75 0.88 0.98 0.99 1.01 1.03 1.24 2 27 2 87 4 10 4 10 4 28 4 12 4 57	1.50 1.66 1.67 1.67 1.77 1.62 1.64	6.59 6.97 6.67 6.76 7.16 6.44 6.10	3.00 2.94 2.70 2.79 2.93 2.70 2.51	1.39 1.30 1.17 1.22 1.27 1.24 1.18	3.56 3.76 3.30 3.47 3.72 3.72	2.43 2.29 2.11 2.10 2.23 2.45 2.57	1.73 1.63 1.50 1.48 1.58 1.75 1.96	1.32 1.24 1.13 1.11 1.19 1.33 1.50		0.58 0.85 0.87 0.87 0.94 1.15	2.61 3.27 3.98 4.12 4.09 4.50 5.73	1.50 1.76 1.81 1.91 1.94 2.52	0.57 0.63 0.76 0.77 0.78 0.80 0.98	2.57 3.00 3.20 3.18 3.33 3.17 3.49	1.15 1.29 1.30 1.29 1.37 1.22 1.21	5.07 5.30 5.03 5.08 5.40 4.73 4.38	2.20 2.14 1.94 2.00 2.10 1.39 1.70	4.37 3.80 7.78 8.16 8.44 8.01 7.32 6	2 00 1 00 1 00 1 10 10 10 10 10 10 10 10	1.09 1.05 0.98 0.93 1.03 1.09 1.01	7.09 5.85 6.37 5.24 5.62 7.18 7.73

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DIURNAL AND ZONAL MEAN OF MID-MONTH VALUES

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